

4.2 Take differences of the two layer eqns. (1)

• define $\Delta u = u_1 - u_2$, $\Delta v = v_1 - v_2$

$$\Rightarrow \Delta u_t - f \Delta v = g' E_x \quad \text{eq. from } \boxed{x \text{ mom}}_1 - \boxed{x \text{ mom}}_2$$

$$\Delta v_t + f \Delta u = g' E_y$$

and forming $H_2 \boxed{\text{mass}}_1 - H_1 \boxed{\text{mass}}_2$

$$\Rightarrow H_2 \left(\frac{1}{\mu} - 1 \right) E_t - H_1 E_t + H_1 H_2 (\Delta u_x + \Delta v_y) = 0$$

recall $E = \mu \eta + \eta = \frac{1}{\mu} E$

The mode with big internal displacements has $[E] \gg [\eta]$

$$\text{so } [\mu] \gg 1 \text{ and } \frac{1}{[\mu]} \ll 1$$

(called the "rigid lid approximation")

$$\Rightarrow \boxed{\text{mass}} \text{ becomes } -E_t + \frac{H_1 H_2}{H_1 + H_2} (\Delta u_x + \Delta v_y) = 0$$

and $\boxed{x \text{ mom}}$ $\Delta u_t - f \Delta v = -g' (-E)_x$

$\boxed{y \text{ mom}}$ $\Delta v_t + f \Delta u = -g' (-E)_y$

These are mathematically identical in form to the Poincaré wave solutions; with

<u>Poincaré</u>		<u>2-Layer</u>	
η	\longrightarrow	$-E$	
u	\longrightarrow	Δu	
v	\longrightarrow	Δv	
H	\longrightarrow	$H_1 H_2 / (H_1 + H_2) \equiv H_{eff}$	the "effective depth"
g	\longrightarrow	g'	

Thus the solutions will satisfy

$$E_{tt} + f^2 E - g' H_{eff} (E_{xx} + E_{yy}) = 0$$

$$\Rightarrow \text{dispersion relation } \omega = (K^2 c^2 + f^2)^{1/2}, \quad K = k^2 + l^2$$

$$\text{and } c = \sqrt{g' H_{eff}}$$

This is the baroclinic mode: zero net transport

$$\text{so eg. } H_1 u_1 + H_2 u_2 = 0, \quad \text{so } H_1 \boxed{x \text{ mm}}_1 + H_2 \boxed{x \text{ mm}}_2 =$$

$$\Rightarrow 0 = -g(H_1 + H_2) \eta_x - g' H_2 E_x$$

or, using $E = \mu \eta$

(3)

$$\Theta = [-g(H_1 + H_2) - g'H_2\mu] \eta_x$$

start here.

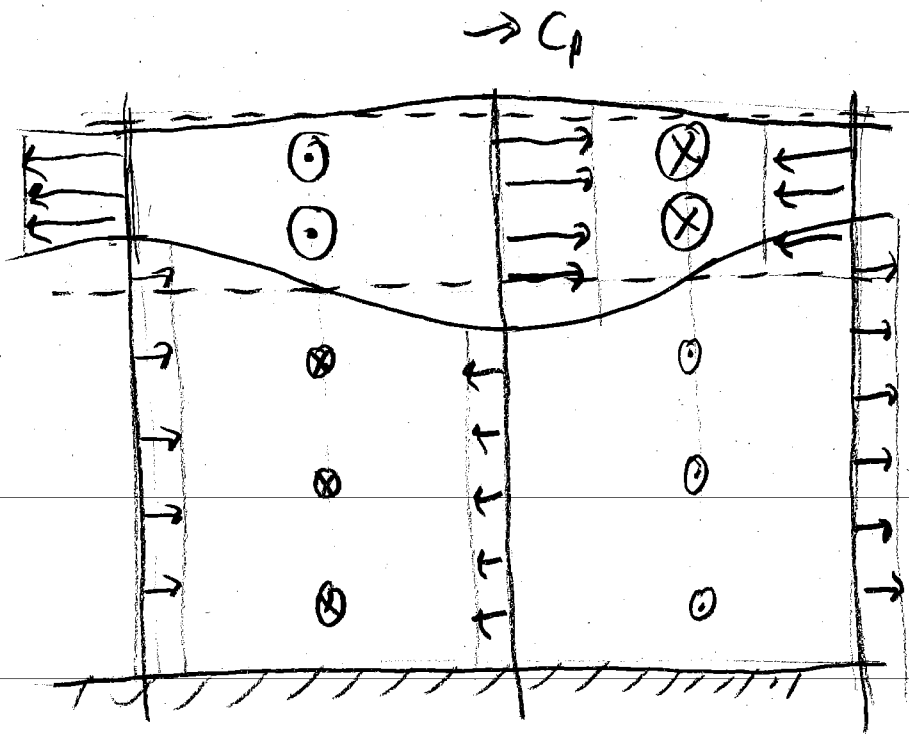
for this to be true for all η_x we require

$$\mu = -\frac{g(H_1 + H_2)}{g'H_2}$$

: - large in magnitude
- negative in sign

} consistent with our earlier assumptions ✓

Solution sketch:



- PV is conserved in each layer
- Ω reverses w/depth
- Ω can be much bigger now because E is big.

The other, "barotropic," mode has $[\mu] \sim \mathcal{O}(1)$ and is almost the same as 1 layer (unstratified) SWE flow.

Barotropic mode	Baroclinic mode
- 1 layer	2 layer
max wave speed = \sqrt{gH}	= $\sqrt{g' H_{eff}}$ slow!
no vertical shear	big vertical shear ($\Delta u, \Delta v$)
Interface follows free surface w/ smaller displacement	Interface <u>down</u> when surface is <u>up</u> , and it moves a lot
$E = \left(\frac{H_2}{H_1 + H_2}\right) \eta$	$E = -\frac{g}{g'} \left(\frac{H_1 + H_2}{H_2}\right) \eta$

eg. ocean thermocline $g' = \frac{\rho \Delta \rho}{\rho_0} = \frac{10 \times 2}{10^3} \frac{m}{s^2} = 2 \times 10^{-2} m s^{-2}$

$H_{eff} = \frac{500 m \cdot 3500 m}{4000 m} = 437 m$

$\Rightarrow C = \sqrt{g' H_{eff}} = 3 m s^{-1}$, vs. $\sqrt{gH} = 200 m s^{-1}$

for atm weather case: $g' = 1 m s^{-2}$, $H_{eff} = 2.5 km$, $C = \sqrt{g' H_{eff}} = 50 m s^{-1}$