

VII.1

12/10/2009

①

COMPRESSIBILITY

< KC 1.8-10 >

- Dry air can be treated as an ideal gas, with equation of state $p = p(p, T)$ derived from

$$p = \rho R T$$

$$R = 287 \text{ J kg}^{-1} \text{ K}^{-1}$$

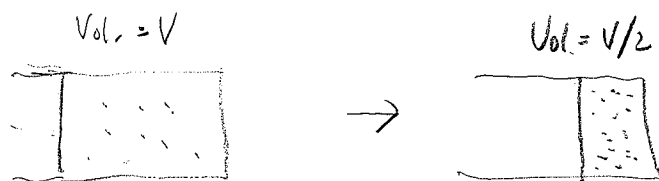
Physical interpretation: the pressure on a wall depends on

(i) how many molecules hit it per unit time and what their mass is $\Rightarrow p \propto \rho$

(ii) the average speed of the molecules $\Rightarrow p \propto T$

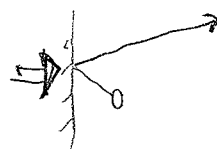


Consider compressing the air in an insulated piston



- Volume decreased $\Rightarrow \rho$ increased $\rho \rightarrow 2\rho$
- T also increased due to KE added to molecules by collision with moving wall

$\therefore p$ more than doubles



Consideration of ideal gas thermodynamics for changes which are

- adiabatic \rightarrow no external loss or gain of heat, and
- reversible (eg. no viscous dissipation)

then the system is "isentropic", and we may show

$$\frac{p}{\rho^\gamma} = \text{const.} \quad \text{where } \gamma = 1.4 \text{ for dry air}$$

or $p = (\text{const.}) \rho^{1.4}$ and still $p = \rho R T$

Now the effect of increasing T is incorporated into the exponent on ρ : $\gamma > 1$

(KC 16.1-2)

Sound waves rely on compressibility for their restoring force, instead of gravity

Assume small variations of $p = \rho$

$$p = p_0 + p'(x,t) \quad [p'] \ll p_0$$

$$\rho = \rho_0 + \rho'(x,t) \quad [\rho'] \ll \rho_0$$

also $[u] \ll c$ the wave speed

Linearize the equations:

x mom $\frac{D u}{D t} = -\frac{1}{\rho} \frac{\partial p}{\partial x} \Rightarrow \frac{\partial u}{\partial t} = -\frac{1}{\rho_0} \frac{\partial p}{\partial x} = -\frac{1}{\rho_0} \left(\frac{\partial p'}{\partial p} \right) \frac{\partial p}{\partial x}$ (1)

mass $\frac{D \rho}{D t} + \rho \nabla \cdot \underline{u} \Rightarrow \frac{\partial \rho}{\partial t} + \rho_0 u_x = 0$ (2)

then forming $\frac{d(1)}{dt} + \frac{d(2)}{dx} \Rightarrow u_{tt} - \left(\frac{\partial p'}{\partial p} \right) u_{xx} = 0$ A wave eqn with phase speed $c_p = \left(\frac{\partial p'}{\partial p} \right)^{1/2}$ (non-dispersive)

Then using $\frac{p}{\rho^\gamma} = \frac{p_0}{\rho_0^\gamma} \Rightarrow p = \frac{p_0}{\rho_0^\gamma} \rho^\gamma \Rightarrow \frac{\partial p}{\partial \rho} = \frac{\gamma p_0}{\rho_0^\gamma} \rho^{\gamma-1} = \frac{\gamma p_0}{\rho_0^\gamma} \rho_0^{\gamma-1} = \gamma \frac{p_0}{\rho_0}$

+ using $p = \rho R T \Rightarrow \frac{\partial p}{\partial \rho} = c^2 = \gamma R T$ $c \approx 340 \text{ m s}^{-1}$ for air at 20°C $c \approx 1400 \text{ m s}^{-1}$ for water

Basic properties of our favorite fluids

Air

 N_2 78.1%, O_2 21.0%, ...

$$p = p(T, p, q) \leftarrow \text{Eqn. of state} \rightarrow p = p(s, T, p)$$

$$\text{specific humidity} = \frac{\text{mass of water vapor}}{\text{mass of moist air}}$$

$$\rho \approx 1.2 \text{ kg m}^{-3} \text{ sea level}$$

$$\rho \rightarrow 0 \text{ high up}$$

Air often approximated as an "ideal" or perfect gas, with

$$p = \rho R T \leftarrow \text{deg. K}$$

$$\text{gas const.} = 287 \frac{\text{J}}{\text{kg K}} \text{ for dry air}$$

Sea Water

 H_2O , NaCl ~ 35‰ (ppt)

$$\text{salinity} \approx \frac{\text{mass of salt}}{\text{mass of seawater}} \times 1000$$

$$\rho = 1000 \text{ kg m}^{-3} \text{ fresh}$$

$$1025 \text{ ocean surface}$$

$$1055 \text{ abyss}$$

change mostly due to compressibility