

William S. Kessler (NOAA/PMEL) and Lionel Gourdeau (IRD-Nouméa)

The problem turned out to be both too easy and too hard ...

- Too easy, because the meridional scale of the ENSO signal is much larger than the Solomons and easily floods around them.
- → The answer is obvious: The Solomons pose no barrier to ENSO wave propagation.
- Too hard, because diagnosing the balance of transport into the Solomon Sea is not presently possible, either from data or models (need ~1/20° resolution).
- → Choose an "easier" piece of the puzzle

What is the structure of flow across 9.5°S into the Solomon Sea?





High salinity traces the thermocline-level flow in the S. Pacific

A tongue subducts in the southeast and flows to the Solomon Sea



The high-S tongue enters the Solomon Sea from the open Pacific Salinity on $\sigma_{\vartheta}=24.5$ CARS data. Overlay geostrophic streamlines 5°S 35.8 55-15 High-S tongue 35.75 35.7 10°S 35.65 35.65 35.62 **Bifurcation** 35.6 of NQC/EAC 15°S 35.55 35 35.5 35.45 20°S 35.4 145°E 150°E 155°E 160°E 165°E

Fewer hints on the intermediate-depth flows ...

But this circulation is crucial for understanding the sources of the ITF, in light of the hypothesis that South Pacific intermediate water is transformed into the shallower, warmer outflow of the ITF.

OGCM meridional current at 10°S with and without an ITF:



Lee et al (2002) MIT OGCM



Mean flow across 9.5°S into the Solomon Sea

ORCA model circulation at surface, thermocline and below



Above 100m: Flow through Sol. St. is <u>southward</u>. (Consistent w/ obs).

Thermocline level:

Sol. St. flow is northward (Pacific inflow $\sim 1/2$ total).





Salinity and Oxygen on Sigma-theta = 27

CARS data. Overlay streamlines on isopycnal



At sigma 27 (~6-800m), the sparse available data suggests that a low-S, high- O_2 tongue penetrates out of the Solomon Sea into the equatorial Pacific via the Australian WBCs.

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that transform SW Pacific intermediate water to the shallower, warmer water that exits into the Indian Ocean.



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Extra Figures Follow

Solomon Island sills



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Overlay ORCA sills



In an OGCM, the NQC and NGCC are uncorrelated at interannual timescales, due to the lagged arrival of ENSO Rossby waves.



Variability of both is well accounted-for by the arrival of linear Rossby waves (mostly due to central Pacific ENSO curl).

There is a factor of more than 2 between the Rossby speed at the NQC (13°S) vs. NGCC (8°S). → ~1-2 year lag

This mostly reflects the n=1 variability.

Rossby model: c=3.0 m s⁻¹, 24-month damping