

# Astrobiological Stoichiometry

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## Abstract

Chemical composition affects virtually all aspects of astrobiology, from stellar astrophysics to molecular biology. We present a synopsis of the research results presented at the “Stellar Stoichiometry” Workshop Without Walls hosted at Arizona State University April 11–12, 2013, under the auspices of the NASA Astrobiology Institute. The results focus on the measurement of chemical abundances and the effects of composition on processes from stellar to planetary scales. Of particular interest were the scientific connections between processes in these normally disparate fields. Measuring the abundances of elements in stars and giant and terrestrial planets poses substantial difficulties in technique and interpretation. One of the motivations for this conference was the fact that determinations of the abundance of a given element in a single star by different groups can differ by more than their quoted errors. The problems affecting the reliability of abundance estimations and their inherent limitations are discussed. When these problems are taken into consideration, self-consistent surveys of stellar abundances show that there is still substantial variation (factors of  $\sim 2$ ) in the ratios of common elements (*e.g.*, C, O, Na, Al, Mg, Si, Ca) important in rock-forming minerals, atmospheres, and biology. We consider how abundance variations arise through injection of supernova nucleosynthesis products into star-forming material and through photoevaporation of protoplanetary disks. The effects of composition on stellar evolution are substantial, and coupled with planetary atmosphere models can result in predicted habitable zone extents that vary by many tens of percent. Variations in the bulk composition of planets can affect rates of radiogenic heating and substantially change the mineralogy of planetary interiors, affecting properties such as convection and energy transport. Key Words: Supernovae—Chemical evolution—Extrasolar planets—Geochemistry—Habitability. *Astrobiology* 14, 603–626.

## 1. Introduction

ASK AN ASTRONOMER to draw the periodic table of the elements, and it may look something like Fig. 1. Stars are almost entirely hydrogen and helium. Using the abundances of Lodders (2010) for the Sun’s initial composition, H and He make up  $\sim 98.5\%$  of the mass. The species O, C, Fe, Ne, Si, N, Mg, and S (in that order) make up  $\sim 1.5\%$  of the mass. However, these elements are central in the thoughts of planetary scientists, geochemists, geophysicists, and astrobiologists as they try to understand how planets form and function and how they evolve as abodes for life. Understanding their origin and distribution are therefore fundamental to un-

derstanding the potential habitability of extrasolar planetary systems.

Whereas H and He are primordial, created in the Big Bang, other elements, referred to as “metals” by astronomers, largely are not. Following the Big Bang, Li, Be, and B represented one-billionth of the mass, and heavier elements were present at a level of one part in  $10^{12}$  (Steigman, 2012). Virtually all the metals present today were created by nucleosynthesis in a variety of stellar environments.

Because Fe is abundant and has many atomic transitions easily observed via optical absorption in a stellar atmosphere, astronomers measure Fe abundance as an indicator of the chemical composition of a star. The Fe/H abundance

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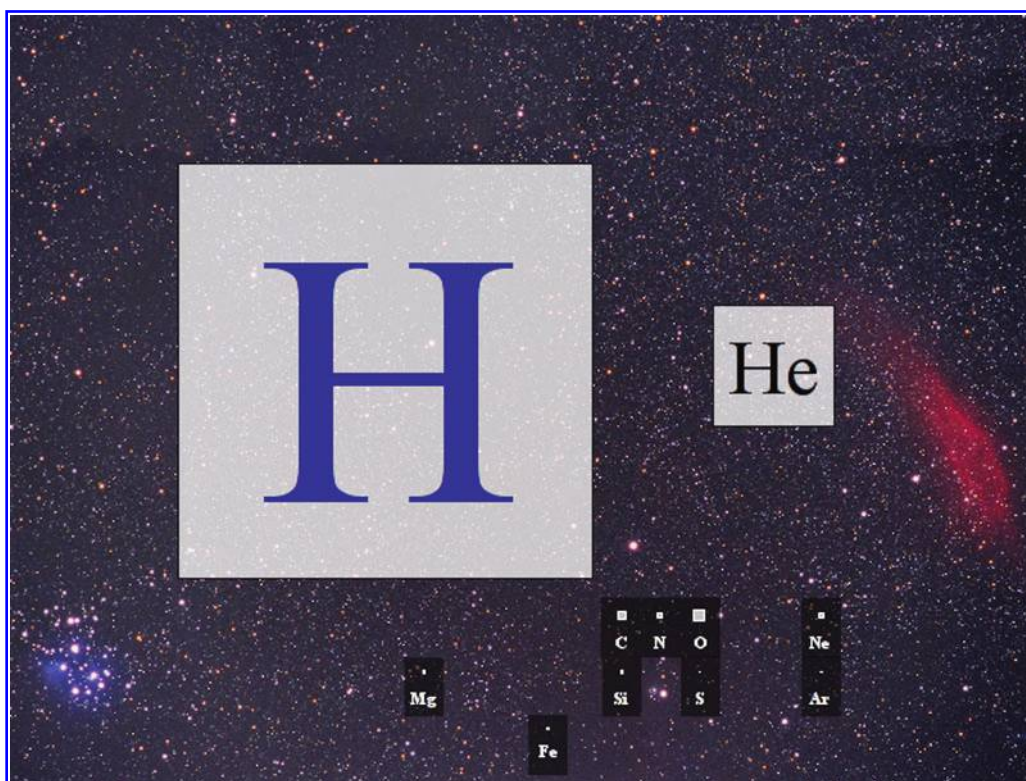
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**FIG. 1.** An astronomer's periodic table of the elements. The size of each element's white square measures the abundance of that element (by number) in the Sun. Elements heavier than He are termed "metals." Because the Sun is considered to sample the average interstellar medium, it is presumed that all stars will share the Sun's relative proportions of metals. The relative sizes of the boxes for C, O, etc., are presumed to be the same for all stars. One of the surprises of stellar spectroscopy studies is that this presumption is false. The observed variations in stellar abundance ratios have significant consequences for star and planet formation evolution and for habitability. (From <http://bjm.scs.illinois.edu/astrochem.php>) Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)

ratio in a star is usually reported as the logarithm of the number ratio of these elements scaled to the ratio in the Sun, denoted  $[\text{Fe}/\text{H}]$  (Helfer *et al.*, 1959). In other words,  $[\text{Fe}/\text{H}] = \log_{10}[(\text{Fe}/\text{H})_{\text{star}}/(\text{Fe}/\text{H})_{\text{Sun}}]$ .

We can discuss other elemental ratios normalized to the Sun in the same way. It is common to use the same notation to quantify elemental ratios relative to Fe, since Fe is most commonly measured. For example,  $[\text{O}/\text{Fe}] = \log_{10}[(\text{O}/\text{Fe})_{\text{star}}/(\text{O}/\text{Fe})_{\text{Sun}}]$ .

The term "metallicity" is commonly used by astronomers. Formally, it refers to the total amount of all heavy elements in the star. Practically, however, metallicity and  $[\text{Fe}/\text{H}]$  are often used interchangeably because thousands of stars have measured  $[\text{Fe}/\text{H}]$ , but many fewer have complete measurements for even the most abundant subset of elements. Using metallicity and Fe content in this way implicitly makes the assumption that *all* metals scale with Fe in the same proportions we observe in the Sun, as if the relative sizes of the boxes in the astronomer's periodic table looked exactly the same as Fig. 1, for *every* star. In this article we will use metallicity to refer to  $[\text{Fe}/\text{H}]$  only. For example, when we say a star has one-third the metallicity of the Sun,  $Z = 1/3 Z_{\text{Sun}}$ , we mean only  $[\text{Fe}/\text{H}] = -0.5$ .

One of the lessons of stellar spectroscopy over the last decade or so is that the elemental abundances of stars do not all scale directly with the metallicity. Data exhibiting variation in abundance ratios has existed for some time, but much improved statistics and recognition of the potential signifi-

cance of this variation were substantially driven by the advent of large-scale efforts at planet detection and characterization. At the same time, a bonanza of exoplanet discoveries has revealed an incredible diversity of planet types of which our solar system represents only a small subsample. The range in elemental abundances and the diversity of planet types are probably not unrelated; these differences can have significant consequences for the evolution of stars, the formation of planets, the processes that occur within planets, and ultimately the habitability of a stellar system.

Following the thread of elements—from astronomical scales to planetary settings to biological ecosystems—has been a major thrust of the Arizona State University "Follow the Elements" team of the NASA Astrobiology Institute (NAI). These themes are sure to figure prominently in future astrobiology research as well. With these thoughts in mind, on April 11–12, 2013, Arizona State University hosted an NAI Workshop Without Walls on the theme of elemental abundance ratios, entitled "Stellar Stoichiometry." The Web site for the conference <http://stellarstoich.sese.asu.edu> (or <https://sites.google.com/a/asu.edu/stellarstoich>) describes the conference and archives electronic versions of the presentations. A companion article (Desch *et al.*, 2014) describes the conference itself. This article represents a synopsis of the scientific presentations from that workshop.

The themes explored in the workshop and in this article are as follows (as this is in part a conference report, the

primary contributors are identified for each section): In Section 2, we discuss the process and results of observationally determining the abundances of stars and planets. Inferring stellar abundances is a challenge, but the data show that elemental abundance ratios do vary significantly from star to star. A link between stellar abundances and planet formation is suggested. We have tantalizing direct measurements of abundances in exoplanets, but the interpretation of those preliminary measurements is far from clear. We also present cautionary tales of attempts to measure abundances of planets in our own solar system. In Section 3, we examine the consequences of elemental abundance variations for aspects of habitability. We consider stellar evolution models and planetary climate models and their effects on the time evolution of habitable zones (HZs). We also consider the consequences for heating and structure and dynamics within planets. In Section 4, we discuss factors that can lead to variable abundance ratios in stars, such as injection of supernova material, and in planets, such as volatile transport in protoplanetary disks. We summarize these diverse findings in Section 5, where we conclude that known elemental abundance ratio variations will have significant consequences for star and planet formation and evolution and for the habitability of a stellar system.

## 2. Finding Abundances of Stars and Planets

The first step in assessing the effect of abundances on any aspect of habitability is to determine the actual abundances of the stars and planets. This is by no means a trivial exercise, even for the stars, which are much brighter and easier to observe. The high-resolution spectra that are needed are burdensome to obtain, and for many nearby stars only the abundances of Fe, or at most a few elements, are determined. The measurement is fraught with uncertainty, often from unrecognized sources. Nevertheless, self-consistent surveys reveal that there are substantial intrinsic variations in elemental abundance ratios, even between stars of similar metallicity.

As difficult as it is to infer stellar abundances, observationally constraining planetary abundances is even more difficult, although it is becoming increasingly tractable. Absorption spectra of gas giant planet atmospheres allow some constraints on elemental abundances, but inferences about bulk abundances in both terrestrial and gas giant planets remain difficult even in the planets of our own solar system.

### 2.1. Stellar abundance measurements (Natalie Hinkel)

Hinkel (2012) compiled spectroscopic abundance data for 50 elements across 3069 stars within 150 pc of the Sun from 84 literature sources to produce the Hypatia Catalog. The goal of Hypatia is to examine nearby stars and regions in the solar neighborhood for chemical abundance trends.

When analyzing the abundance measurements taken across 25 years from different groups at a variety of instruments, there is potential for many variations between the data sets. While some effort has been made to compare new abundance determinations with more standard benchmark catalogs, for example, Edvardsson *et al.* (1993), Thévenin and Idiart (1999), or Valenti and Fischer (2005), those comparisons are not standardized, varying in many ways from the statistical to graphical. The unfortunate result is that different groups often measure significantly different

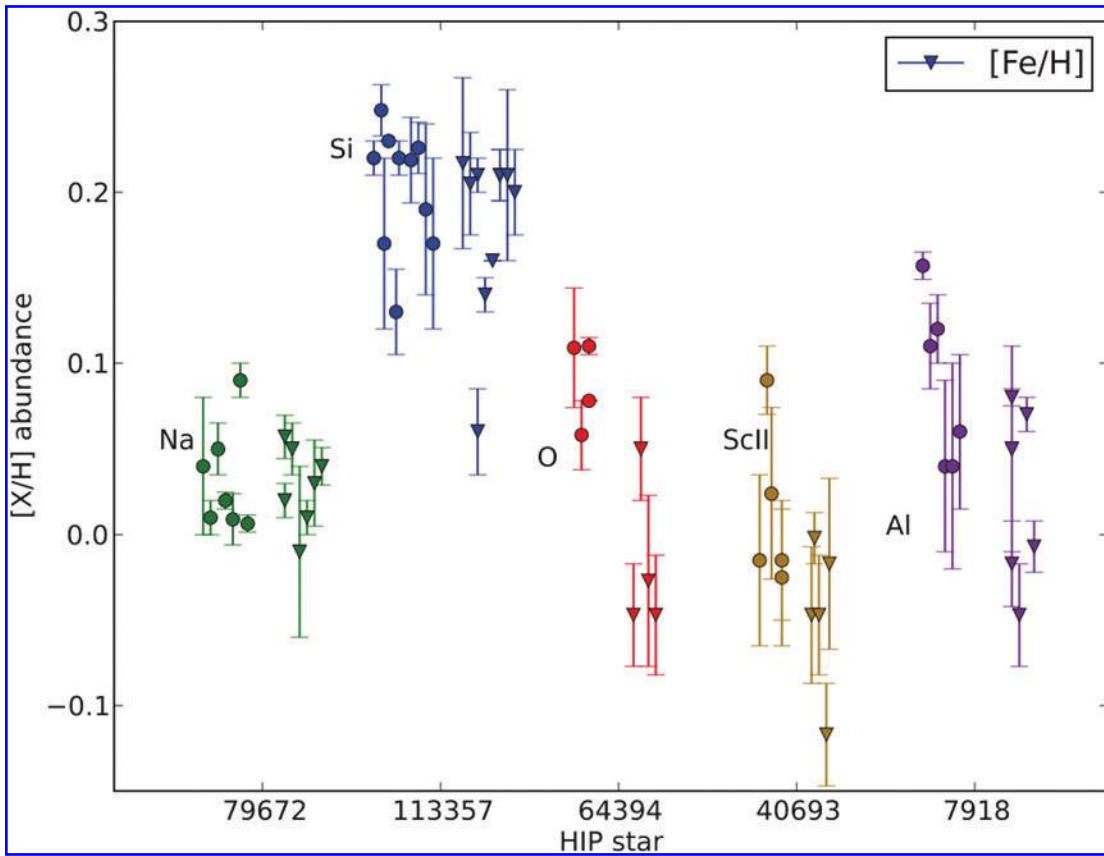
abundances for the same element in the same star. Figure 2, from the work of Hinkel (2012), illustrates how different research groups measured the elements Na, Si, O, Sc, Al, and Fe in five different stars (the data sets are discussed and the research groups identified by Hinkel, 2012). The maximum abundance measurement minus the minimum for each element within a star is referred to as the “spread.” The observational spread in [Na/H] within HIP 79672, for example, is  $\sim 0.1$  dex, which greatly exceeds the quoted observational errors for many groups,  $\sim 0.04$  dex. Different groups must be measuring stellar abundances using inconsistent methodologies, since the variation in abundance determinations goes far beyond observational errors.

One complication is that different groups scale to different assumed solar abundances such as those of Holweger (1979), Anders and Grevesse (1989), and Asplund *et al.* (2009), although ultimately this is not the main source of the discrepancy. The differences between solar abundance determinations are dishearteningly large. The quoted abundances for the primordial solar composition are derived from a mixture of techniques, primarily analysis of primitive meteorites and spectroscopy of the solar photosphere (Lodders, 2010). Photospheric abundances face two problems. First, the measured abundances must be corrected for stellar processes such as gravitational settling of heavy elements over the life of the Sun by using stellar models. The second and larger problem is that the translation of measured widths of spectral lines to an actual number of atoms requires a stellar atmosphere model. The 3-D models of the solar photosphere and convection zone that have been developed over the last decade give substantially different abundances for some species than the old 1-D approximations (Asplund *et al.*, 2009). One of the most strongly affected species is oxygen, the third most abundant element in the Universe.

All the abundances within the Hypatia Catalog were corrected for the solar normalizations applied by each individual group and then renormalized according to the same scale, namely, Lodders (2010) as described by Hinkel and Kane (2013). For example, Valenti and Fischer (2005) determined that HIP 490 has [Ti/H]=0.05, using the solar abundance scale of Anders and Grevesse (1989) for which  $\log_{10}\epsilon(\text{Ti})=4.99$ . (Since solar abundances are absolute, we cannot use the [X/H] notation. In this notation,  $\log_{10}\epsilon(X)$  is the total number of atoms of X normalized to  $\log_{10}\epsilon(\text{H})=12.00$ ). For comparison, Lodders (2010) gave  $\log_{10}\epsilon(\text{Ti})=4.93$ . Renormalizing the Valenti and Fischer (2005) result [Ti/H]=0.05 to the Lodders (2010) abundance scale, we find [Ti/H]= $0.05+4.99-4.93=0.11$  dex. When this method was applied to all the elements within the Hypatia stars, the median absolute difference between the original and renormalized abundances was 0.04 dex. The change resulting from renormalization was often larger than the reported error, but the total effect of applying a consistent solar renormalization on the abundance spread was minimal, actually increasing it from 0.14 to 0.15 dex.

It is clear from the analysis performed on the Hypatia Catalog that not only are there some major discrepancies between different stellar abundance data sets for the same elements in the same stars but the remedy must be more than cosmetic. An evaluation of the stellar parameters, such as effective temperature and surface gravity, reported by multiple authors indicates that stellar atmosphere models





**FIG. 2.** Abundances of various elements in five stars, as measured by different groups, taken from various catalogs (Hinkel, 2012). For HIP 79672, measurements of Na (circles) and Fe (triangles) are displayed. For HIP 113357, measurements of Si (circles) and Fe (triangles) are displayed, etc. Error bars denote uncertainties quoted in the cited catalogs. The cumulative spread (maximum abundance determination minus the minimum) is larger than the typical quoted errors of the individual measurements. Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)

might be the source of the problem. However, the varying number and quality of spectral lines might also cause the large amount of spread. Even atomic data such as oscillator strength can be shown to cause substantial variations. Many of these factors are examined systematically in a companion paper arising from the workshop and in preparation for the *Astrophysical Journal* (Hinkel *et al.*, unpublished). This work is receiving contributions from five independent groups applying their methods to the same set of observational data. Fixed values for stellar parameters, atomic data, equivalent width measurements of spectral line strength, and line lists are provided so that each factor may be independently varied and the source of discrepancies identified between methods. It is hoped that this will result in a recommended set of best practices.

The preliminary results for four stars and three groups presented at the workshop itself indicate that line lists do have a substantial effect. Differences in temperature  $T$ , surface gravity  $\log(g)$ , and  $[\text{Fe}/\text{H}]$  were up to 266 K, 0.05 dex, and 0.36 dex, respectively. Imposing fixed stellar parameters typically changed  $\Delta[\text{Fe}/\text{H}]$  between groups by 0–0.08 dex. Most other elements exhibited the same behavior. Large disparities exist for elements that have either one weak line that is difficult to measure or have many lines of which a subset are chosen for doing the analysis. In at least one case a difference in the oscillator strength used for

calculation of electron transitions in Mg produced a significant disparity.

## 2.2. Variations in stellar abundance ratios of nearby stars (Michael Pagano)

To judge the physical effects of elemental abundance ratios, we must find statistical methods to separate the observational and methodological sources of scatter from the physical variation in abundance ratios exhibited by stars, which we refer to as the “intrinsic scatter.” Disentangling the intrinsic scatter from the actual observational errors is not straightforward. To do so, we performed a linear regression ( $\chi^2$  fit) (Isobe *et al.*, 1990) in which the total scatter is assumed to be the sum in quadrature of an observational uncertainty  $u_{\text{obs}}$  and an intrinsic scatter  $u_{\text{intrinsic}}$ . The goodness-of-fit parameter  $\chi^2_\nu$  of the observed points is

$$\chi^2_\nu = \frac{1}{n} \sum_{i=1}^n \frac{(X_i - \mu_i)^2}{u_{\text{obs}}^2 + u_{\text{intrinsic}}^2}$$

where  $n$  is the number of measured stars,  $X_i$  is the measured abundance ratio in star  $i$ ,  $\mu_i$  is the average of all  $X_i$ , and again  $u_{\text{obs}}$  is the observational uncertainty and  $u_{\text{intrinsic}}$  represents the real, physical scatter in the abundance ratios. (*A priori* we do not expect a normal distribution, so we do not use  $\sigma$  to

represent the scatter in the denominator.) If there is a real intrinsic scatter in the stars, correctly estimated errors should not be able to account for all the observed dispersion. For each element and for each star, there is a quoted observational error  $u_{\text{obs}}$ , so all variables except  $u_{\text{intrinsic}}$  are fixed. We vary the intrinsic scatter term until  $\chi^2_{\nu} = 1$  (De Silva *et al.*, 2006; Stuart and Ord, 2009). If  $\chi^2_{\nu} > 1$  without a nonzero  $u_{\text{intrinsic}}$  term, this would mean there is more dispersion in the data than can be accounted for by observational errors alone. For  $\chi^2_{\nu} \leq 1$ , the spread for that element might be entirely explained by the observational error, although an intrinsic variation may yet exist at a small value hidden in generously estimated observational uncertainties. The inclusion of the intrinsic uncertainty term in this formulation also allows us to avoid unwanted emphasis on a few well-measured points.

Table 1 summarizes the calculated values of  $u_{\text{intrinsic}}$  for six different abundance studies (Bond *et al.*, 2006, 2008; Ramírez *et al.*, 2007; Takeda, 2007; Mishenina *et al.*, 2008; Neves *et al.*, 2009; González Hernández *et al.*, 2010). Of these studies, only Bond *et al.* (2006, 2008), González Hernández *et al.* (2010), Ramírez *et al.* (2007), and Neves *et al.* (2009) gave errors for individual stars. The other data sets gave the average error per element for the entire study. This lessens the robustness of this statistical method, as  $u_{\text{intrinsic}}$  directly links observational error to spread on a star-to-star basis; use of an average error still allows for the calculation and comparison of  $u_{\text{intrinsic}}$  but only provides a lower limit. Studies with larger observational errors, such as that of Delgado Mena *et al.* (2010), tend to not exhibit a  $u_{\text{intrinsic}}$  value.

The relative pattern of  $u_{\text{intrinsic}}$  between elements is consistent across all studies that yield a useful  $u_{\text{intrinsic}}$ . For example, C, Na, Si, Ca, and Cr have lower  $u_{\text{intrinsic}}$  than other elements below the Fe peak, in all cases. Few studies

look at elements above the Fe peak produced by the  $r$ - and  $s$ -processes, but those that do find that heavier elements tend to have a large spread. Likewise, all groups tend to find a larger spread for O, Mg, and Al. This seems to indicate that these elements have the largest actual scatter, but it must be noted that the larger  $u_{\text{intrinsic}}$  for oxygen may be attributable to the difficulty in measuring O absorption lines and to a discrepancy in measurements due to the 6300 Å forbidden line (Nissen, 2013). This line is strongly affected by a blend with a line of Ni I with strength  $\sim 55\%$  of the O I line in the Sun. At higher metallicities or lower abundances of O (relative to Fe-peak elements), the contribution can be even larger. The [O I] line at 6363 Å also has a small contribution from a CN molecular line in cool stars (Caffau *et al.*, 2013) but is generally considered more reliable than the 6300 Å line. The O I triplet at 7774 Å is not affected by blends but is sensitive to non-local thermodynamic equilibrium effects (Takeda, 2003; Ramírez *et al.*, 2007; Fabbian *et al.*, 2009). From our intrinsic variation analysis, we expect reliable results only from the 6363 Å line and 7774 Å triplet. To the extent that groups include the 6300 Å line in their analysis, this may explain some of the apparently intrinsic scatter. The values quoted here refer to studies that do not use the 6300 Å line.

Despite these caveats, the most striking fact about these abundances is that the light elements do show significant intrinsic variation. There are many stars that seem to have compositions significantly different from solar, by about a factor of 2 from the average in each direction.

### 2.3. Planetary abundances (Nikku Madhusudhan)

2.3.1. Abundances of gas giant exoplanets. Recently, it has become possible to estimate elemental abundance ratios in the atmospheres of “hot Jupiters,” or highly irradiated

TABLE 1. INTRINSIC VARIATIONS OF ELEMENTAL ABUNDANCE RATIOS FOR DIFFERENT ABUNDANCE STUDIES

Element	Intrinsic Variation					
	Bond06&08	GH10	Mish08	Neves09	Ramírez07	Take07
C	0.017	0.045				
O	0.092	0.052	0.153		0.155(L) and 0.197(N)	
Na	0.036	0.034	0.013	0.052		0
Mg	0.068	0.029	0.116	0.054		0.095
Al	0.047	0.044	0.112	0.097		0.085
Si	0.014	0.014	0.062	0.040		0.050
S		0.180				0
Ca	0.016	0.007	0.065	0.028		0
Sc		0.025	0	0.090		0.094
Ti	0.026	0.002	0.061	0.083		0.080
V		0	0.037	0.151		0
Cr	0.019	0	0.063	0		0
Mn		0.057		0.051		0
Co		0.031	0.080	0.063		0.028
Ni	0	0.020	0	0.017		0
Cu		0				0
Zn		0.035	0.034			0
Sr		0.000				
Y	0.032	0.033				
Zr	0.113	0				
Ba	0.117	0.087				
Ce		0.059				
Nd	0.033	0				
Eu	0.135	0.01				

gas giant exoplanets. Significantly, the O/H and C/O ratios are easier to measure for atmospheres of hot Jupiters than they are for Solar System giant planets. The vast majority of known extrasolar gas giants have  $T \sim 600\text{--}3000$  K, thus hosting gaseous  $\text{H}_2\text{O}$  in their upper atmospheres that is observable in their spectra. Other detectable gases include  $\text{CH}_4$ ,  $\text{CO}$ ,  $\text{CO}_2$ , and  $\text{NH}_3$ , depending on the temperature and incident irradiation (Burrows and Sharp, 1999; Lodders and Fegley, 2002; Madhusudhan, 2012). Measurements of these molecular abundances in the  $\text{H}_2$ -rich atmospheres of hot Jupiters allow estimations of elemental abundance ratios involving H, C, O, and N. While a statistically significant constraint on the C/O ratio of an exoplanetary atmosphere has been possible only recently (Madhusudhan *et al.*, 2011a), nominal constraints on atmospheric C, O, and/or C/O have been reported for several transiting exoplanets in the past decade (Fortney *et al.*, 2005; Seager *et al.*, 2005; Desert *et al.*, 2009; Madhusudhan and Seager, 2009; Swain *et al.*, 2009; Madhusudhan, 2012).

Beginning with the first inference of  $\text{C/O} > 1$  for WASP-12b (Madhusudhan *et al.*, 2011a), several other hot Jupiters are now suggested as possibly C-rich (Madhusudhan, 2012), though such inferences are still very preliminary given the low resolution of the data. Nevertheless, high C/O ratios present a radical departure from our understanding of solar system chemistry. Most stars are oxygen-rich, with  $\text{C/O} < 1$  (Delgado Mena *et al.*, 2010; Petigura and Marcy, 2011), like the Sun ( $\text{C/O} = 0.5$ ; Asplund *et al.*, 2009). Assuming similar O-rich chemistry for gas giant planets, their nebular compositions would imply that  $\text{H}_2\text{O}$  is the dominant molecule in their atmospheres and easily observable for hot Jupiters. At temperatures common to hot Jupiters ( $T > 1500$  K), and at C/O ratios slightly higher than solar, the molecular chemistry in  $\text{H}_2$ -rich atmospheres varies significantly from the familiar Solar System chemistry (Madhusudhan *et al.*, 2011b; Kopparapu *et al.*, 2012; Madhusudhan, 2012; Moses *et al.*, 2013). For example, at  $T = 2000$  K, as the C/O ratio is increased from the solar value of 0.5 to a C-rich value of 1.1, the  $\text{H}_2\text{O}$  abundance drops by 2 orders of magnitude, and the  $\text{CH}_4$  abundance increases by over 3 orders of magnitude. Moreover, other hydrocarbons such as  $\text{C}_2\text{H}_2$  and  $\text{HCN}$  also become markedly abundant for  $\text{C/O} > 1$ , whereas  $\text{CO}_2$  becomes negligible. Such extreme changes in observable chemistry of hot Jupiters as a function of C/O ratio make them ideal laboratories for constraining the elemental abundance ratios of C/H, O/H, and C/O.

Estimates of high C/O ratios in giant exoplanets raise the question of how a C-rich gas giant could form around an O-rich star, motivating new ideas on the formation of planetary systems. Lodders (2004) suggested the possibility that Jupiter, which the Galileo probe measured to have  $\text{C/O} > 1$  (at least locally: see Subsection 2.4), could have formed by accreting tar-dominated planetesimals instead of the water ice-dominated planetesimals expected in the Solar System based on the composition of minor bodies. Lodders (2004) suggested a high abundance of such planetesimals just beyond a “tar line” in the solar nebula (see also Subsection 3.2). In a similar vein, following the more recent inference of  $\text{C/O} > 1$  in WASP-12b, Öberg *et al.* (2011) suggested that C/O ratios in giant planetary envelopes depend on the formation location of the planets in the disk relative to the condensation fronts, or ice lines, of major C- and O-bearing volatile species such as  $\text{H}_2\text{O}$ ,  $\text{CO}$ , and  $\text{CO}_2$ , since the C/O

ratio of the gas approaches 1 outside the CO and  $\text{CO}_2$  ice lines. By predominantly accreting such C-rich gas, more so than O-rich planetesimals, gas giants could, in principle, host C-rich atmospheres even when orbiting O-rich stars. On the other hand, it may also be possible that inherent inhomogeneities in the C/O ratios of the disk itself (*e.g.*, by local concentrations of carbonaceous grains) may contribute to higher C/O ratios of the planets relative to the host stars (Kuchner and Seager, 2005; Madhusudhan *et al.*, 2011b; Mousis *et al.*, 2012). Internal processes of the planets themselves may cause the composition of the observable parts of the atmosphere to inaccurately reflect the bulk composition (see also Subsection 2.4.1). These scenarios predict different limits on the metallicities and C/O ratios of the giant planets, which high-precision observations of their atmospheres will be able to test in the near future but for which knowing the abundances of the host stars is a central requirement.

**2.3.2. Abundances of rocky exoplanets.** In contrast to gas giant planets, rocky planets form almost entirely from refractory condensates, possibly with some volatile condensates, but with very little gas accreted from the disk. Rocky planet compositions are governed by what minerals condense near the planet’s location in a protoplanetary disk with distinct elemental composition. Bulk compositions for terrestrial-type planets are expected to span a wide range based on models (*i.e.*, Carter-Bond *et al.*, 2012) and on observations of our own solar system. The potential for forming systems with significantly nonsolar ratios is significant. Wide ranges of C/O and Mg/Si values have been observed in the photospheric abundances of exoplanet host stars. The frequency of very C-rich stars ( $\text{C/O} > 0.8$ ) is being debated, but some examples certainly exist, and large Mg/Si variations are in little doubt (Delgado Mena *et al.*, 2010; Petigura and Marcy, 2011; Fortney, 2012; Nissen, 2013).

The advent of high-precision measurements of masses and radii of transiting super-Earths (exoplanets with radii between 1 and 10 Earth radii) are making it feasible to place nominal constraints on the range of their possible interior compositions. Where available, the abundances of the host stars can significantly improve the constraints on the super-Earth compositions by constraining the mineralogy. Masses and radii are known for  $\sim 15$  transiting super-Earths to date. Models of super-Earths used to interpret masses and radii in the past have traditionally assumed Earth-like oxygen-rich interiors made of Fe, silicates, and water (Valencia *et al.*, 2007, 2010; Charbonneau *et al.*, 2009; Rogers and Seager, 2010; Winn *et al.*, 2011; Gillon *et al.*, 2012). Recently the first inference of a C-rich super-Earth was reported for 55 Cancri e (Madhusudhan *et al.*, 2012). The star was known to be carbon-rich (Delgado Mena *et al.*, 2010), but the C/O ratio of the system may be lower than originally thought (Teske *et al.*, 2013). Using the planet’s mass and radius, Madhusudhan *et al.* suggested a carbon-rich interior in the planet, with carbon (as graphite and diamond), iron, silicon carbide, and/or silicates as the primary constituents and water being scarce. The drastically different thermochemical properties of carbon-rich minerals open an entirely new regime of geological processes (see Subsection 3.4).

## 2.4. Lessons from our solar system

**2.4.1. Inferring gas giant planet compositions** (Steven Desch). Since the detection of Na in the atmosphere of HD

209458b (Charbonneau *et al.*, 2002), transmission spectroscopy has offered the promise of constraining the abundances of gas giant exoplanets. Important to note, however, is that such measurements probe only the upper atmospheres of gas giant planets and not necessarily their bulk abundances. The difficulties of translating observable abundances into bulk abundances are perfectly illustrated by the example of Jupiter in our own Solar System.

The Galileo entry probe's Neutral Mass Spectrometer measured the abundances of H, He, C, N, O, Ne, S, Ar, Kr, and Xe in Jupiter's atmosphere, between the 0.88 and 21.7 bar levels (Niemann *et al.*, 1998; Mahaffy *et al.*, 2000; Wong *et al.*, 2004). The abundance of P was measured remotely from the Galileo orbiter's Near Infrared Mapping Spectrometer (Irwin *et al.*, 1998). In Fig. 3, we plot the elemental ratios He/H, C/H, and so on, normalized to the values in the solar nebula (Asplund *et al.*, 2009). The species C, N, S, P and the noble gases Ar, Kr, and Xe are all enriched relative to solar by a similar factor of 3. He, in contrast, is depleted by 15% below the protosolar value, and Ne is depleted by a factor of 6. O is depleted by a factor of 2.

The enrichment of most species is potentially explained by external photoevaporation of the Sun's protoplanetary disk by far-ultraviolet (FUV) radiation from a nearby massive star

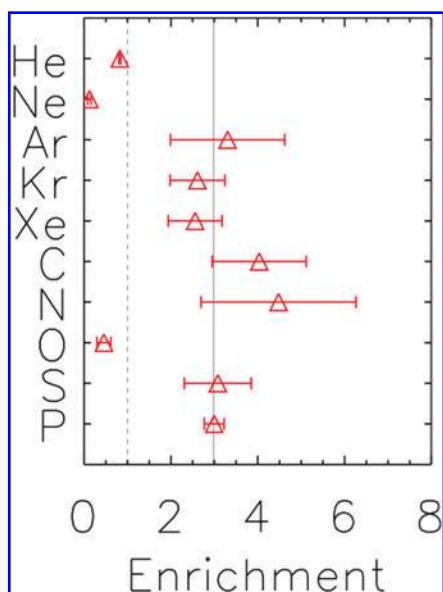
(Guillot and Hueso, 2006). Within a thin ( $<0.1 \text{ g cm}^{-2}$ ) layer, gas is heated to temperatures  $\approx 500\text{--}1000 \text{ K}$ , sufficient to drive the gas to escape the gravitational pull of the star. If the FUV radiation is propagating into gas that normally is at temperatures  $<25 \text{ K}$ , it is possible that only  $\text{H}_2$ , He, and Ne remain in the gas phase, as even  $\text{N}_2$  and Ar can be trapped in amorphous water ice at these cold temperatures (Bar-Nun *et al.*, 1985; Owen *et al.*, 1992). Although Guillot and Hueso (2006) only considered photoevaporation off the top and bottom surfaces of the disk, where temperatures probably are not  $<25 \text{ K}$ , Adams *et al.* (2004) showed that when considering a full disk, most mass loss will take place at an outer edge of the disk,  $r_d$ . For typical FUV fluxes ( $G_0=3000$  in Habings,  $1.6 \times 10^{-3} \text{ erg cm}^{-2} \text{ s}^{-1}$ , equivalent to an O6 star at 2 pc) and typical protoplanetary disk parameters,  $r_d=50 \text{ AU}$  (Adams *et al.*, 2004). (This is also the location of the edge of the Solar System's Kuiper belt: Morbidelli *et al.*, 2003.) At this distance in a protoplanetary disk, temperatures are cold enough that all species except H, He, and Ne are likely to be trapped in icy bodies and *not* removed with the photoevaporated gas. Jupiter easily could have accreted from gas that was not so much enriched in the heavier species but depleted in H.

Loss of  $\text{H}_2$ , He, and Ne from a photoevaporated disk explains why other species are enriched but does not explain the depletions in He and Ne. The leading explanation for these depletions is demixing at high pressures. At depths of several thousand kilometers and pressures of several megabar,  $\text{H}_2$  is predicted to transition from a molecular gas to metallic H, separating out from He in the process. He droplets must form that sink because they are denser than the surrounding fluid. Precipitation of He to Jupiter's core over many billions of years is thought to be the cause of Jupiter's He depletion (Stevenson and Salpeter, 1977a, 1977b). It is hypothesized that Ne would partition into the He droplets but Ar would not, a result qualitatively confirmed by Wilson and Militzer (2010) using density functional theory–molecular dynamics calculations.

This leaves only the case of O as anomalous. Soon after the Galileo results showed that other species were enriched but O was depleted, it was suggested that the probe sampled a water-poor region of Jupiter. This explanation is very plausible given that the probe entered one of Jupiter's "5  $\mu\text{m}$  hotspots," downwellings of gas analogous to high-pressure cells, from which water has been removed (Atreya *et al.*, 1999). The Juno mission, arriving at Jupiter in 2016, will globally probe the O/H ratio down to the 200 bar level to test this hypothesis (Matousek, 2007).

An alternative is that Jupiter accreted gas with an O/H ratio 3 times solar, and then depleted by a factor of 6 similar to Ne, ending up with a factor of 2 depletion. Dissolution of Ne but not Ar into He droplets is controlled to a large degree by atomic radius (Wilson and Militzer, 2010). N, C, Ar, Kr, and Xe have atomic radii larger than that of H, while He and Ne have atomic radii smaller than H (Clementi *et al.*, 1967). The atomic radius of O is not as small as Ne, but it is smaller than H and may be expected to dissolve in the He droplets as well. If Juno measures a global depletion in O/H, it will be difficult to interpret the result in any way except that O, like He and Ne, is sequestered in Jupiter's core.

$\text{H}_2\text{O}$  and  $\text{CH}_4$  are some of the species most easily measured by transmission spectroscopy, and the C/O ratio that can be derived from these measurements is crucial to



**FIG. 3.** Abundances of elements in the upper atmosphere of Jupiter, as measured by the Galileo entry probe and orbiter. Abundances are reported relative to H abundance and relative to solar composition (Asplund *et al.*, 2009). C, N, S, P, and even Ar, Kr, and Xe are apparently uniformly enriched by a factor of 3 above solar. He is depleted, presumably because of H/He demixing at megabar pressures. Ne is depleted by a factor of 6, presumably because of sequestration into precipitated He droplets. Galileo measured a factor of 2 depletion in O, but it was argued to be a local anomaly. Alternatively, this may be evidence that O was enriched by a factor of 3 like other species, then depleted by a factor of 6 like Ne, due also to sequestration into precipitated He droplets. Surface measurements of O may not reflect bulk abundances. Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)



chemical models of planet formation. But there are no guarantees that measurements of a gas giant planet's upper atmosphere will reflect its bulk abundance or the star's abundance. If O can be sequestered in a gas giant planet, transmission spectroscopy of its upper atmosphere will overestimate the C/O ratio.

2.4.2. Inferring terrestrial planet compositions (Miriam Riner). Beginning just in 2009, it has become possible to measure the mass and radius of exoplanets not much more massive than Earth, such as CoRoT-7b (Queloz *et al.*, 2009) and GJ 1214b (Charbonneau *et al.*, 2009 and subsequent papers). With the mass and radius known, it is possible to estimate the density and thereby infer the bulk composition. In practice, what are published are mass-radius relationships (*e.g.*, Fortney *et al.*, 2007; Seager *et al.*, 2007; Barnes *et al.*, 2013). In these plots, individual curves of radius versus mass are computed for an assumed composition (or at most two compositions, one in a core and one in a mantle). The mass-radius error ellipses of exoplanets are plotted over these lines to infer their compositions. Here "composition" refers to the relative abundance of iron metal, rock, and water ice. Through use of such plots, it has been inferred that Kepler-11b is likely to be more water ice than silicate rock, and CoRoT-7b is more likely to be more Fe metal than silicate rock (Kaltenegger *et al.*, 2013).

These models of mass-radius relationships build on those developed for studying planets in our own solar system, inheriting their insights *and* their built-in assumptions. We must understand how simplifying assumptions and uncertain input parameters translate into uncertainties in the mass-radius relationships and inferred compositions.

Mass-radius relationships must be computed because materials under planetary interior pressures change density. This is done by using laboratory-determined properties of the materials assumed for a metallic core, rocky mantle, and/or icy layer, together with an adopted equation of state (EOS), to compute the effects of self-compression on the interiors of model planets. Different approaches vary in detail based on the adopted EOS and other assumptions. As an example, Riner *et al.* (2008) modeled the self-compression of the Solar System terrestrial planets, especially Mercury, using the Birch-Murnaghan finite strain EOS and assuming an adiabatic temperature profile. Planets are divided into multiple layers within which material properties are assumed to be uniform. For each material or layer, the model requires the density of the material at standard temperature and pressure ( $\rho$ ), the bulk modulus ( $K$ ), the pressure derivative of the bulk modulus ( $K'$ ), the thermal expansion coefficient ( $\alpha$ ), and the second Grüneisen parameter ( $\delta$ , which describes the relationship between bulk modulus and density). By using these material properties, together with an assumed temperature difference across thermal boundary layers where the composition changes, the effect of self-compression on the volume of a planet of a given composition can be computed, leading to mass-radius relationships.

Because a composition is assumed, then a mass-radius relationship calculated, there are several limits to using such relationships to infer internal composition; a single mass-radius pair does not uniquely constrain the composition of a planet. The mass and radius of Mercury have been accurately known for six decades (Urey, 1951), yet its metal-to-

silicate ratio (or the percent of mass in the core) was poorly known until new spacecraft measurements of its moment of inertia. Until recently, published estimates of the core mass fraction ranged anywhere from 50% to 100% of the total planetary mass (Harder and Schubert, 2001; Chen *et al.*, 2007; Riner *et al.*, 2008).

Much of this uncertainty has stemmed from poor knowledge of the composition (abundance and identity of a light-alloying element) and phase (liquid or solid) of the core. For example, liquid Fe is 20% less dense and 40% more compressible than solid iron (face-centered cubic phase). While these differences tend to diminish at higher pressures, uncertainties in material properties also are greater at higher pressures. To a lesser degree, uncertainties in mantle composition also matter. The major mantle compositional variation relevant to mass-radius relationships in our Solar System is the relative abundance of Fe and Mg in the mantle (the Mg#). While the Mg# makes a substantial ( $\sim 40\%$ ) difference in density, the compressibility and other material properties vary less than those due to compositional variations in core materials. That said, the uncertainty in the mass-radius relationship due to variations in Mg# has not been documented over a wide range of planetary masses or Mg# variations.

Additional uncertainty stems from the difficulty of measuring EOS data. The material properties of candidate compositions at planetary interior pressures and temperatures must be determined through static diamond-anvil cell experiments (for pressures up to  $\sim 100$  GPa, equivalent to Earth's core mantle boundary) or shock-wave experiments (for pressures up to  $\sim 300$  GPa, Earth's inner core but temperature is not controlled). Mass-radius relationships can only be constructed assuming those materials that have been measured in these ways.

The level of uncertainty about the composition and structure of our own planet, and the degree to which other sources of information are used, is often underappreciated by scientists working outside the field in astrophysics or planetary science. From seismic measurements, we know that the density of the molten outer core of Earth requires the presence of one or more light-alloying elements. Usually this element is assumed to be S, in part because data are lacking for other candidate alloys (H or Si in FeNi). Although composing only a few percent of the core's mass, this alloying element can make a substantial difference in the material properties of the core and therefore the assumed bulk composition of the planet. In practice, we use cosmochemical abundances measured in meteorites and the solar atmosphere to make reasonable assumptions about the identity of the light alloying element.

The assumptions we typically make for our Solar System may not be valid in other stellar systems. Sulfur can be redistributed during planet formation, and rocky exoplanets may have different S abundances in their cores, affecting our interpretation of their bulk rock/metal ratio. The observed variations in stellar Mg/Si ratios, if acquired by these stars' planets, may introduce high-pressure mineral phases in their mantles that do not exist in Earth's (as discussed in Subsection 3.4).

To make substantial progress in deriving accurate mass-radius relationships for exoplanets, it will be necessary to experimentally measure a greater range of material



properties spanning the potential compositional variability. At the same time, the uncertainties in input parameters and the limitations of mass-radius relationship modeling should be understood, especially by astronomers, to avoid overly interpreting the observational data.

### 3. Effects of Composition on Stars and Planets

Even though virtually all stars are overwhelmingly (>98% by mass) H and He, a star's evolution will be significantly affected by the abundances of those trace (<2%) elements heavier than He. Modeling the effect of composition on stars is essential but is a *relatively* simple problem compared to planets. The structure and internal processes of a planet are, of course, completely changed depending on its composition: whether it is a gas giant, Neptune-sized ice giant, or a rocky planet, or simply subtle variations of these classes.

As can be expected, the habitability of a planet is even more sensitive to the composition of the planet itself. The ability to keep liquid water on an Earth-like planet's surface (the usual criterion of habitability) depends on the atmosphere of the planet, which in turn depends on its composition. There are also composition-dependent processes internal to an Earth-like planet that can affect its surface, such as radiogenic heating, internal mineralogy and rheology, thermal evolution, and convective mixing.

#### 3.1. Abundance variations and stellar and habitable zone evolution (Patrick Young)

Models of the evolution of stars of different metallicity are generally created assuming the heavy elements are found in the same proportions relative to Fe as in the Sun. Except for uniform enhancement of the  $\alpha$  elements (O, Ne, Mg, Si, S, Ar, Ca, Ti) in stars of very low metallicities, variations in abundance ratios at a given [Fe/H] are neglected, despite the fact that stellar evolution is sensitive to the star's composition.

The sensitivity to composition arises from two effects: the EOS and the radiation opacity (Iglesias and Rogers, 1996; Rogers *et al.*, 1996). The changes in the EOS are relatively minor, but rearranging the proportions of different species at a constant [Fe/H] can result in opacity changes of tens of percent (Iglesias and Rogers, 1996). Opacity changes the rate of leakage of radiation, and increased radiation pressure in the stellar envelope drives expansion, resulting in larger radii and lower effective temperatures. A slower rate of energy loss also requires a slower rate of nuclear burning to maintain hydrostatic equilibrium. So we expect stars with enhanced abundances to be cooler, of lower luminosity, and longer-lived relative to stars with the same [Fe/H].

The large observed intrinsic spread in [O/Fe] (see Subsection 2.2) and the high abundance of O make this element particularly interesting. In stellar interiors, it is second only to Fe in contributing to the opacity. Stellar evolution simulations [herein using the TYCHO code (Young and Arnett, 2005)] with variable abundance ratios show that the impact is dramatic (Young *et al.*, 2012).

In these models, a solar composition from the work of Lodders (2010) was adjusted to the mean abundance ratios of the sample in Subsection 2.2, with all other elements being maintained at solar values for the standard composition. The intrinsic physical spreads  $u_{\text{intrinsic}}$  in the element-

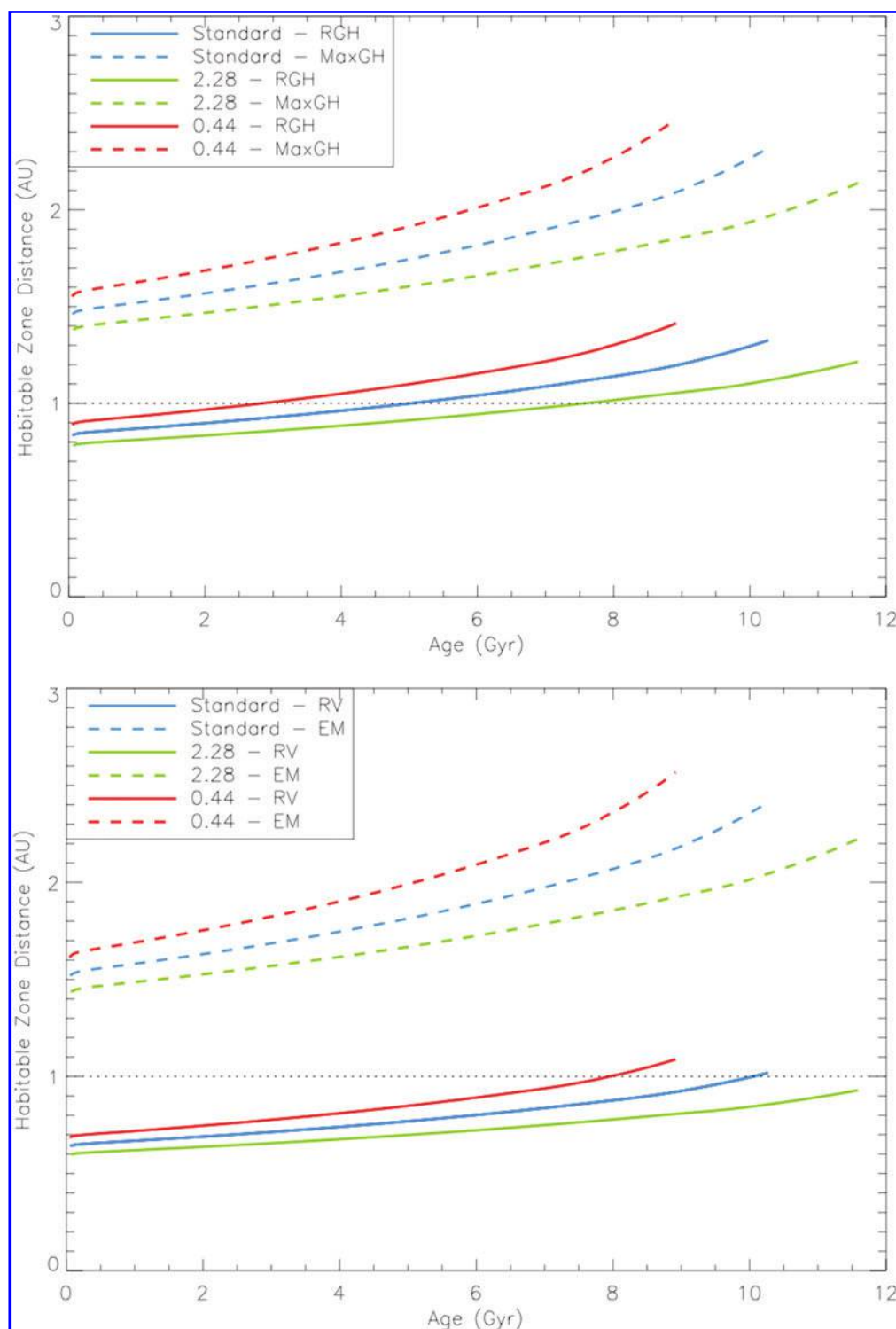
to-Fe ratios are the basis for the variant models. In each model, one element was varied to the  $3u_{\text{intrinsic}}$  extremes of the distribution, resulting in one standard composition model and two extreme models for each of the elements C, O, Na, Mg, Al, Si, Ca, Ti. Custom opacity tables for each composition were obtained from the OPAL project (Iglesias and Rogers, 1996). All models were run until core hydrogen exhaustion. The models provide high time-resolution histories of luminosity and effective temperature.

The prescription of Kopparapu *et al.* (2013), a conceptual descendant of Kasting *et al.* (1993), was used to convert modeled stellar effective temperature  $T_{\text{eff}}$  and luminosity versus time into HZ boundaries. This treatment allows the HZ edges to be easily scaled from the stellar quantities for different assumptions, and the relative differences between the stellar models are robust.

In all cases, enhanced abundance ratios cause larger changes than depletions. C, Na, and Mg have small but noticeable effects. C has a high abundance but relatively few electron transitions and low ionization potential. Mg and Na have lower abundances but higher opacity per gram than C, resulting in a similar degree of shift in the tracks. Si has less impact due to its smaller range of variation, and Al, Ca, and Ti have very small effects due to their small abundances; these can be neglected for the purpose of habitability. The largest changes, unsurprisingly, arise from variation in O. Both the luminosity and  $T_{\text{eff}}$  of enriched compositions are systematically lower at a given age, but the most profound effect is on the pace of the evolution. The main-sequence turnoff for the low-[O/Fe] model occurs at an age of 9 Gyr. The standard composition (which is slightly more O-rich than solar) has a turnoff age of 10.5 Gyr, and the O-rich model turns off at 12 Gyr (and at lower luminosity and  $T_{\text{eff}}$ ).

This time evolution of the HZ is highly significant. If the ultimate goal is the discovery of life on an Earth-like world, one must consider not just the current habitability of the planet but also how long the planet has remained in its HZ. Life did not modify Earth's atmospheric composition at levels that might be detectable by near-future astronomical observations until about a billion years after its formation. Complex life on Earth developed during the late Ediacaran period, 635–542 Myr ago, some 4 Gyr after its formation. Orbits with semimajor axes near the edges of HZs will not remain there for the entire lifetime of the star. Figure 4 shows the inner and outer edges of HZs for high, low, and standard [O/Fe] as a function of time, calculated by using conservative and optimistic cases chosen from the parameterization of Kopparapu *et al.* (2013). The dashed line marks a 1 AU orbit. The conservative case (top) uses the prescriptions for the Runaway and Maximum Greenhouse models. The optimistic case (bottom) uses the Recent Venus and Early Mars values.

For a star otherwise identical to the Sun in mass and Fe abundance at the lower end of the range of [O/Fe] observed in planet host candidates, an Earth-like planet at 1 AU from the star would not (given the assumptions about planetary atmospheres herein) remain habitable long enough for complex life to have developed at the pace that occurred on Earth. The effect is obviously dramatic for the extremes of O variation, but the effect could be important to less extreme systems with planets on the edges of their nominal HZs. For a set of stars deliberately chosen to be maximally Sun-like by Ramírez *et al.* (2009), the smaller range of [O/Fe] still causes the 1 AU habitability



**FIG. 4.** Inner and outer edges of the HZ versus time for variations in O/Fe for conservative (top) and optimistic (bottom) HZ predictions. Habitable zones are shown for models with depleted (upper; red), standard (middle; blue), and enhanced (lower; green) O abundance ratios. Inner edges are solid; outer edges are dashed. A 1 AU orbit is indicated by the dashed horizontal line. The conservative case uses the prescriptions for the Runaway (RGH) and Maximum (MaxGH) Greenhouse cases of Kopparapu *et al.* (2013). The optimistic case uses the recent Venus (RV) and early Mars (EM) values. Depletions systematically move the HZs outward due primarily to increased luminosity and secondarily to increased  $T_{\text{eff}}$ . The difference in location increases markedly with age. The most dramatic effect results from the substantial difference in main sequence lifetimes. Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)

lifetime to differ by nearly a billion years. The magnitudes of the effect for the observed ranges of  $[C/Fe]$ ,  $[Na/Fe]$ , and  $[Mg/Fe]$  are hundreds of millions to a billion years.

Only a very limited range of orbits remain continuously habitable. For the conservative atmospheric prescription, there are no orbits continually habitable for all the models. Even the most optimistic case considered has orbits that are continuously habitable for all compositions extending only from 1.1 to 1.4 AU. Ranges are similarly narrow, though shifted in radius, for other  $[Fe/H]$ . The range of orbits that provide at least 4 Gyr of continuous habitability are much more extensive. High- $[O/Fe]$  models have the widest range of 4 Gyr habitable orbits by a small margin due to their longer overall main sequence lifetime and shallower rate of change of the zone boundaries. A point that is not commonly considered is that the majority of orbits that are at some point habitable only become so during the latter part of the star's lifetime. While these abundance variations encompass the majority of nearby planet host candidates, there are even more extreme outlier systems. The stars HD 53705 ( $[O/Fe] = 0.6$ ) and HD 136352 ( $[C/Fe]$  also  $\sim 0.6$ ) have a substantially larger range of orbits that remain habitable for a multi-billion year window due to their large enhancements.

The stellar characteristics that determine the size and persistence of HZs change at an important level with different abundance ratios. Therefore proper characterization of a system requires more data than just  $[Fe/H]$ .

### 3.2. Habitable zones from atmospheric models and the frequency of potential habitable planets in extrasolar planetary systems (Ravi Kopparapu)

The HZ is classically defined as the circumstellar region in which a terrestrial-mass planet ( $0.1 \leq M_p \leq 10 M_{\text{Earth}}$ ) with favorable atmospheric conditions can sustain liquid water on its surface (Huang, 1959; Hart, 1978; Kasting *et al.*, 1993; Selsis *et al.*, 2007; Kopparapu *et al.*, 2013). Liquid water is often mentioned as a prerequisite for life, and this restriction is appropriate for remote detection of life on planets around other stars. Life as we know it is carbon-based and depends on liquid water.

The average temperature on Earth's surface is approximately 289 K, set by the flux of radiation received from the Sun and also by our atmosphere. On an airless body with the same albedo as Earth (0.3), 1 AU from our sun, the black-body surface temperature would be only 255 K ( $-18^\circ\text{C}$ ), incompatible with liquid water, at least on the surface of the planet. More importantly in that case, of course, liquid water is not stable at any temperature at sufficiently low atmospheric pressure. The atmospheric greenhouse effect, which is attributable mostly to  $\text{H}_2\text{O}$  and  $\text{CO}_2$ , completely changes the habitability of a planet and the HZ boundaries.

Following Kopparapu *et al.* (2013), and assuming an Earth-like planet and atmospheric composition, two kinds of HZ limits are considered:

(1) *Conservative HZ*: The inner edge of the HZ in this limit is fixed by the *moist-greenhouse effect*, which is reached once the surface temperature exceeds  $\sim 340$  K (for a 1 bar Earth-like atmosphere) and the  $\text{H}_2\text{O}$  saturation mixing ratio at the surface exceeds 0.2. This results in a large increase in the altitude of the tropopause and a corresponding increase of water vapor in the stratosphere. For

Earth, the moist-greenhouse limit is 0.99 AU (or in terms of stellar flux normalized to the flux at Earth's orbit,  $S_{\text{eff}} = 1.014$ ) (Kopparapu *et al.*, 2013). If the surface temperature exceeds 647 K (the critical temperature for pure  $\text{H}_2\text{O}$ ), the entire ocean should evaporate. This describes the *runaway greenhouse* limit, which is reached at 0.97 AU ( $S_{\text{eff}} = 1.038$ ) for Earth. The outer edge of the HZ is determined by the *maximum greenhouse* limit, which is the distance at which increasing  $\text{CO}_2$  partial pressures in the atmosphere would no longer result in additional greenhouse warming; increasing  $\text{CO}_2$  partial pressures would tend to lead to surface cooling (Kasting *et al.*, 1993; Kopparapu *et al.*, 2013). For Earth, the *maximum greenhouse* limit is at 1.68 AU ( $S_{\text{eff}} = 0.35$ ).

(2) *Optimistic HZ*: For the inner edge of the HZ in our solar system, an empirical estimate can be calculated based on the observation that Venus could have had significant water up to as late 1 Gyr ago. The Sun at that time was  $\sim 8\%$  less bright than it is today. Therefore, we can derive a *recent Venus* limit for the inner edge, which is at 0.75 AU ( $S_{\text{eff}} = 1.77$ ). Similarly, an empirical estimate for the outer edge of the HZ can also be derived based on the observation that Mars appears to have had liquid water flowing on its surface at or before 3.8 Gyr ago, when the solar flux was some 25% lower (Kopparapu *et al.*, 2013). This *early Mars* limit corresponds to a distance of 1.76 AU ( $S_{\text{eff}} = 0.32$ ).

Parametric forms of both the conservative and optimistic estimates of the HZ for stars of different spectral types were given by Kopparapu *et al.* (2013) and are displayed in Fig. 5. Using these estimates, and six quarters of Kepler data, Kopparapu *et al.* (2013) showed that the occurrence rate of terrestrial-sized ( $0.5\text{--}2 R_{\text{Earth}}$ ) planets around M dwarfs can vary from 48% (conservative estimate) to as high as 61% (optimistic estimate). So the potential for finding Earth-like planets around M dwarfs may be higher than previously reported (Dressing and Charbonneau, 2013).

It should be noted that the HZ estimates discussed above used 1-D radiative-convective climate models that do not include cloud feedback. Results from 3-D global climate models that do include cloud behavior (Abe *et al.*, 2011; Wordsworth *et al.*, 2011; Yang *et al.*, 2013) show that the HZ may be wider; hence the terrestrial planet occurrence rate quoted above may be larger than the previous estimates. This bodes well for future discovery missions like the Transiting Exoplanet Survey Satellite and for future characterization missions such as the James Webb Space Telescope as they attempt to find the first habitable planet outside our Solar System, at least for planets with atmospheric compositions like Earth's.

The distinct possibility exists that an exoplanet may have Earth's mass and orbit 1 AU from a Sun-like star yet not be habitable, because its atmospheric composition is completely different. Earth's atmosphere contains  $4 \times 10^{-4}$  bar of the major greenhouse gas  $\text{CO}_2$  but would contain tens of bars of  $\text{CO}_2$  without the carbonate-silicate cycle. The same assumptions that optimistically place the outer edge of the HZ at 1.76 AU could also involve an inner edge beyond 1.0 AU. Enhanced atmospheric  $\text{CO}_2$  abundance could be due to increased volatile abundance on the planet or to more subtle transport processes. Conversely, in a system with high C/O, C will preferentially be incorporated into other species. A highly reducing environment may be dominated by  $\text{CH}_4$ , a more effective greenhouse gas than  $\text{CO}_2$ .



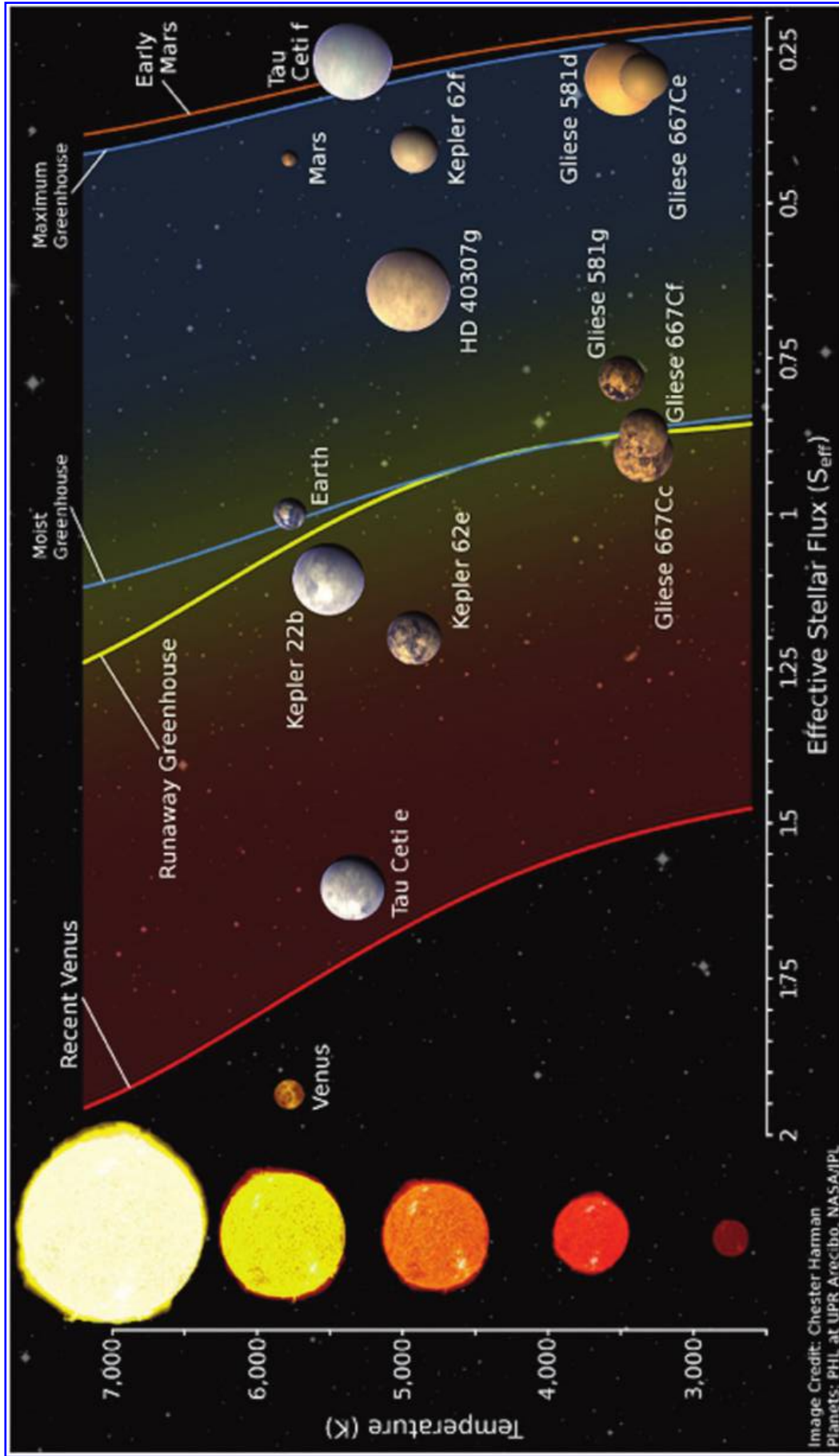


Image Credit: Chester Harman  
Planets: PHL at UPR, Arecibo, NASA/JPL

**FIG. 5.** Incident stellar flux on a planet as a function of stellar effective temperature,  $T_{\text{eff}}$ . The region between the moist and runaway greenhouse cases (cyan and yellow curves) and the maximum greenhouse (blue curve) represents a conservative estimate of the HZ. The optimistic HZ limits are recent Venus (red curve) and early Mars (orange curve). Examples of terrestrial exoplanets are included. Input stellar fluxes with different spectra substantially change the amount of stellar flux needed to produce habitable conditions when realistic energy transport in atmospheres is considered. (Image credit: Chester Harman, Planetary Habitability Lab at UPR, Arecibo, NASA/JPL) (Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast))

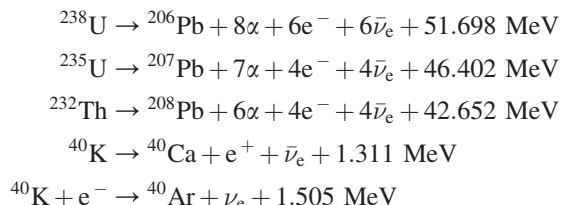
We must consider more than a potentially very different primordial volatile inventory. The surface exchange and mantle recycling processes that modify atmospheres may operate very differently than on Earth if the solid rock geochemistry is exotic. It is difficult to overstate the importance of chemical composition on the contribution of atmospheres to habitability, an area that is largely unexplored at this time.

3.3. Internal heating of terrestrial exoplanets (Rory Barnes)

On Earth, plate tectonics and the hydrologic cycle combine to create the carbonate-silicate cycle that prevents the secular buildup of CO<sub>2</sub> in the atmosphere, staving off a sterilizing moist greenhouse (Walker *et al.*, 1981). Without an internal energy source to sustain plate tectonics, Earth may have become a barren wasteland by today.

Three primary sources of energy are postulated: latent heat from formation, gravitational energy released by differentiation, and radiogenic heating from the decay of radioactive isotopes. Beyond Earth, tidal heating provides significant energy on several icy satellites (see, *e.g.*, Peale *et al.*, 1979; Hurford *et al.*, 2007) and may be important for planets in the HZ of low-mass stars (Jackson *et al.*, 2008; Barnes *et al.*, 2013).

The primary sources of radiogenic energy in Earth are <sup>40</sup>K, U, and Th. Their decay chains are

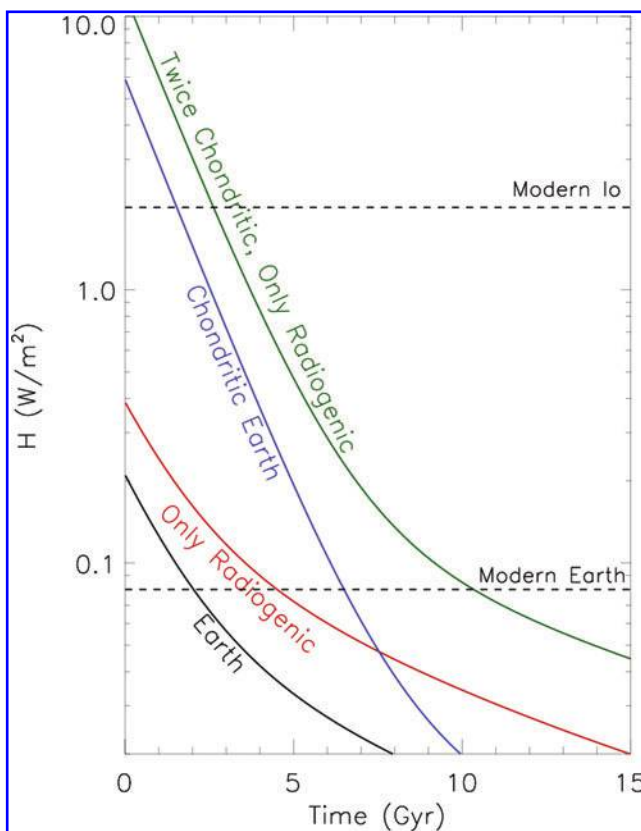


(*e.g.*, Dye, 2012). <sup>40</sup>K decays to <sup>40</sup>Ca 89.3% of the time. The energy produced by these reactions (kinetic energies of alpha and beta particles, plus gamma rays produced by decays and electron-positron annihilations) is the radiogenic heating of Earth. The antineutrinos of the <sup>238</sup>U and <sup>232</sup>Th decay chains escape and do not heat Earth but are detectable by the KamLAND and Borexino neutrino observatories (Araki *et al.*, 2005; Dye, 2012; Ludhova, 2012; Tolich, 2012). These experiments have shown that these isotopes account for >5 TW (probably ~19 TW) of internal energy production, out of a total rate of ~44 TW (Araki *et al.*, 2005; KamLAND Collaboration, 2011; Dye, 2012). For this study, we will assume a Th/U ratio of 3.9 (Anders and Ebihara, 1982; Anders and Grevesse, 1989), similar to chondritic, and a low, but observationally permitted, K/U ratio of 12,000 (Arevalo *et al.*, 2009), ~40 times less than chondritic (Anders and Grevesse, 1989). Earth is enriched in U relative to chondrites but depleted in K. The half-lives of the <sup>235</sup>U, <sup>238</sup>U, <sup>232</sup>Th, and <sup>40</sup>K species are 0.7, 4.5, 13.9, and 1.4 Gyr, respectively. Therefore, K dominates early on, but Th dominates on long timescales.

Figure 6 shows the nominal evolution of Earth’s surface energy flux over time due to radiogenic heating (assuming the energy is instantly transported to the surface), as well as several alternative scenarios. The “Earth” curve is the

above model in which 19 TW (8 TW each from U and Th, 3 TW from <sup>40</sup>K) is generated today. The “Only Radiogenic” curve assumes that the entire 44 TW is generated radiogenically, with the same elemental proportions as the “Earth” case. The “Chondritic Earth” case assumes Earth formed its current U inventory but with 40 times more <sup>40</sup>K than the “Earth” case, and the “Twice Chondritic, Only Radiogenic” case has 80 times the <sup>40</sup>K proportion and twice the U and Th proportion as the “Only Radiogenic” model. These latter curves are plausible for terrestrial exoplanets that formed beyond the snow line and migrated in during the protoplanetary disk phase (Carter-Bond *et al.*, 2012) and hence might not be depleted in K. Note that, if Earth formed with a chondritic K/U ratio, it could have experienced Io-like heat fluxes at its surface for ~1.5 Gyr.

In addition to radiogenic heating, tidal heating could be significant on terrestrial bodies in the HZ of low-mass stars. The tidal heating in the interior is a strong function of composition, as well as external factors such as the mass of the star and the orbit. We consider tidal heating as predicted



**FIG. 6.** Surface energy fluxes due to radiogenic heating, including comparisons to modern Earth and modern Io. Black curve: Our nominal model for Earth in which modern Earth produces 19 TW of radiogenic power coming from 8 TW each of U and Th and 3 TW from <sup>40</sup>K. Red curve: The total 44 TW of internal power of Earth is generated by ~18 TW each of U and Th, and ~7 TW from <sup>40</sup>K. Blue curve: Same as the black curve but with 40 times the K, *i.e.*, chondritic K/U ratio. Green curve: The power from Earth if all its current energy were radiogenic, but then scaled to twice the chondritic K/U ratio. Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)

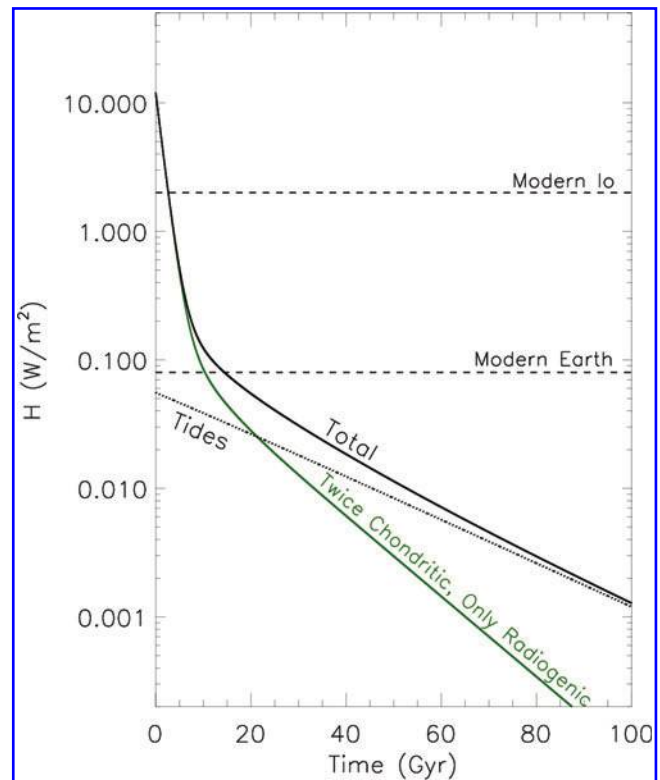
by the “constant-time-lag model” (*e.g.*, Hut, 1981; Ferraz-Mello *et al.*, 2008; Leconte *et al.*, 2010) and compare tidal heating to radiogenic heating on a hypothetical terrestrial exoplanet. For the sake of brevity, we do not reproduce the model here but refer the reader to Appendix E of Barnes *et al.* (2013) for a complete description.

As a specific example, we consider an Earth-mass, Earth-radius planet orbiting Barnard’s Star with an initial eccentricity of 0.19 and a semimajor axis of 0.06 AU so that it experiences approximately the same insolation as Earth. Such a planet would likely have formed at larger distances and migrated in, as accumulation of local material is unlikely to form such a “large” planet around such a small star (Raymond *et al.*, 2007). We assume the planet has a tidal time lag of 640 s, similar to the value for modern Earth (Lambeck, 1977), and allow the planet’s orbit and, hence, tidal heating to evolve with time. We assume 90% of the tidal dissipation occurs in a putative ocean and only 10% in the rocky interior. In addition, we assume the planet has a “Twice Chondritic, Only Radiogenic” composition. Initially, the radiogenic heating dominates, primarily due to  $^{40}\text{K}$ ; then after about 5 Gyr, tides dominate, as shown in Fig. 7. While this planet would experience Io-like conditions for its first 1–2 Gyr, it would subsequently settle into a relatively stable situation during which its internal energy would remain with a factor of 10 of the modern Earth for the next 50 Gyr. As Barnard’s Star is expected to survive for trillions of years, tidal heating could in principle sustain geochemical cycles like plate tectonics and the carbonate-silicate cycle far longer than radiogenic heating, especially if another planetary companion can maintain a nonzero eccentricity. A tidally heated exoplanet with a composition different from Earth’s could easily have different material properties that affect tidal dissipation.

#### 3.4. Internal structure and dynamics of terrestrial exoplanets (Sang-Heon Shim)

While there are differences in their compositions, possibly due to alteration of composition during planet formation and differentiation (Ringwood, 1977; Zindler and Hart, 1986), the abundances of nonvolatile elements in the solar photosphere, type I carbonaceous chondrites (CI chondrites), and Earth’s upper mantle (Taylor and McLennan, 1985) are very similar. This gives hope that stellar elemental abundances may provide us with important constraints on the bulk chemical compositions of planets around stars. We explore two examples that show how variations in some key elemental ratios may affect the composition, structure, and dynamics of planetary interiors. Even for our own planet, many key issues are still under debate, such as the chemical composition of the deep mantle and the core and the phase relations and the physical properties of key constituents (minerals and melts). Therefore, we will limit our discussions to changes in major mineralogy with simplified compositions and their possible impacts on just a few properties of planetary interiors.

**3.4.1. C/O ratio.** Some stars have much higher C/O than that of the Sun (Bond *et al.*, 2010a, 2010b), and the elevated C/O may result in carbide mineralogy, instead of silicate mineralogy, during nebular condensation (Larimer, 1975). Stars with high C/O could host carbide planets (silicon car-



**FIG. 7.** Comparison of tidal and radiogenic heating on a hypothetical Earth-mass, Earth-radius planet in the HZ of Barnard’s Star, initially with a semimajor axis of 0.06 AU and eccentricity of 0.19. The planet’s radiogenic inventory is the same as the “Twice Chondritic, Only Radiogenic” curve from Fig. 6. The tidal heating of the rocky interior is generated by the slowly damping semimajor axis and shown by the dotted curve. The green curve is the same as in Fig. 6. The solid black curve is the total heating flux. Despite the orbital evolution, the planet remains in the insolation HZ for the full 100 Gyr shown. Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)

bides and even elemental carbon in the mantle and iron carbides in the core). Some key thermal and rheological properties of SiC (moissanite) and SiO<sub>2</sub> (quartz) are compared in Table 2. It is important to note that Earth’s mantle consists mostly of magnesium silicates, not pure silica. Furthermore, high pressure in the planet’s interior can change the physical properties of minerals. While extensive studies have been made of the properties of silica and silicates at high pressure–temperature conditions found in planetary interiors (Hemley, 1998, and references therein), much less is known for the properties of SiC at high pressure and temperature. Therefore, the comparison of the properties at 1 bar and 300 K discussed below is to highlight the fundamental differences in physical properties between silicates and carbides.

Silicon carbide polymorphs have very dense crystal structures even at 1 bar, while silica polymorphs have open crystal structures at the same conditions, resulting in fairly large contrast between them at 1 bar. With compression, SiO<sub>2</sub> undergoes a series of polymorphic phase transitions to denser structures: coesite to stishovite at  $\sim 9$  GPa (Stishov and Popova, 1961) and stishovite to seifertite at  $\sim 120$  GPa (Grocholski *et al.*, 2013). The density of seifertite at 1 bar is



TABLE 2. COMPARISON OF SOME PHYSICAL PROPERTIES OF  $\text{SiO}_2$  AND  $\text{SiC}$  AT 1 BAR AND 300 K

Physical properties	$\text{SiO}_2$	$\text{SiC}$
Density ( $\text{g cm}^{-3}$ )	2.66	3.21
Thermal expansion parameter ( $10^{-6} \text{ K}^{-1}$ )	12.3	4.0
Thermal conductivity ( $\text{W m}^{-1} \text{ K}^{-1}$ )	1.3	500

$4.355 \pm 0.003 \text{ g cm}^{-3}$ , (compared to  $2.93 \text{ g cm}^{-3}$  for coesite) demonstrating profound impact of pressure (Grocholski *et al.*, 2013). In contrast,  $\text{SiC}$  does not appear to have a major phase transition up to 90 GPa (Yoshida *et al.*, 1993).

More drastic contrasts can be found as well.  $\text{SiC}$  has a factor of 3 smaller thermal expansivity and a factor of 400 greater thermal conductivity. Of course, it is uncertain whether the magnitudes of these differences will remain the same at higher pressures. Nevertheless, this comparison shows that  $\text{SiC}$  may not gain buoyancy as much as silicates do by heating, and heat will be transported more effectively through conduction in  $\text{SiC}$  than in silicates, suggesting thermal conduction will be much more important as a heat transport process in carbide mantles than in Earth's silicate mantle, where convection is dominant (Davies, 1999). Because surface tectonics are closely related to internal processes, at least for Earth (through mantle convection), the possible dominance of thermal conduction over convection could result in significantly different surface processes in carbide planets, perhaps a lack of plate tectonics and/or lack of mixing between the mantle and surface.

It is important to point out that there remain important questions for the identification of carbide planets. The existing report on the identification of carbide planets (Madhusudhan *et al.*, 2012) relies on knowledge of the thermal equations of state of  $\text{SiC}$  and diamond at high pressures, but the thermal equation of state of  $\text{SiC}$  has been estimated based on low-pressure spectroscopic measurements (Aleksandrov *et al.*, 1989; Seager *et al.*, 2007). Furthermore, the EOS does not take into account a phase transition in  $\text{SiC}$  reported at 90 GPa, which may increase density drastically ( $\sim 10\text{--}20\%$ ) (Yoshida *et al.*, 1993). For reliable identification of carbide planets, it therefore is urgent to establish a robust EOS of  $\text{SiC}$  at high pressure and temperature, including the effects of densification through the phase transition.

Recently, Sen *et al.* (2013) suggested substitution of C for O in  $\text{SiC}$  may be far more energetically possible at high pressures than substitution of C for Si. If this indeed happens at high pressure and forms  $\text{SiO}_x\text{C}_{4-x}$  type materials for carbon-rich planets, the different properties of such materials will have profound impact on internal dynamics and evolution of the planets. Yet it should be further examined if such materials can be actually stabilized under possibly very reducing conditions of carbon-rich planets.

**3.4.2. Mg/Si ratio.** Earth's upper mantle is composed mainly of olivine,  $(\text{Mg,Fe})_2\text{SiO}_4$ , and pyroxene,  $(\text{Mg,Fe,Ca})(\text{Al,Si})\text{O}_3$  with  $\text{Mg/Si} = 1.2\text{--}1.3$  (Ringwood, 1975; McDonough and Sun, 1995). With an increase in depth or pressure, these minerals undergo polymorphic phase transitions. Although this interpretation remains controversial, the  $\text{Mg/Si}$  ratio may remain the same or similar throughout Earth's

mantle, because the boundary between the upper mantle and the lower mantle of Earth can be explained by a polymorphic phase transition, the post-spinel transition (Ringwood, 1970; Ito and Takahashi, 1989; Shim *et al.*, 2001).

Significant variations in  $\text{Mg/Si}$  among different stars have been documented (Bond *et al.*, 2010a, 2010b), ranging from 0.8 to 2 (see also Subsection 2.2), and this variation will result in significant changes in the mineralogy of the mantles in Earth-like exoplanets. For example, compared with Earth's mantle, a decrease in  $\text{Mg/Si}$  will increase the amount of pyroxene and its high-pressure polymorphs and decrease the amount of olivine and its polymorphs. It is even possible to generate free silica ( $\text{SiO}_2$ ), if  $\text{Mg/Si}$  is sufficiently lower than 1, as shown for basaltic compositions (Hirose *et al.*, 2005; Ricolleau *et al.*, 2008). Free silica is not expected to exist in Earth's pyrolytic mantle. On the other hand, an increase in  $\text{Mg/Si}$  to  $\geq 2$  will yield upper mantles almost entirely made of olivine, and it is even possible that ferropericlase,  $(\text{Mg,Fe})\text{O}$ , may exist from very shallow depths. In Earth's mantle, ferropericlase exists only in the lower mantle where olivine breaks down to perovskite,  $(\text{Mg,Fe})(\text{Al,Si})\text{O}_3$ , and ferropericlase,  $(\text{Mg,Fe})\text{O}$ .

This may have profound impacts on mantle dynamics. For example, ferropericlase is known to have much smaller viscosity than other minerals in Earth's lower mantle, which is made of 70%  $(\text{Mg,Fe})(\text{Al,Si})\text{O}_3$  perovskite, 25% ferropericlase, and 5%  $\text{CaSiO}_3$  perovskite. If the amount of ferropericlase increases to a level where ferropericlase grains are mechanically interconnected, the viscosity of the lower mantle may decrease by 2 orders of magnitude (Ammann *et al.*, 2011). Such a significant decrease in viscosity would result in much more vigorous thermal convection in the mantle.

The change in  $\text{Mg/Si}$  will also affect many other processes in the mantles of silicate planets, such as melting relations and phase transitions. For example, melting behavior has a profound impact on planet differentiation. Therefore, early evolution of silicate planets with different  $\text{Mg/Si}$  may result in very different internal structures. Melting behaviors also affect near-surface processes. They will affect formation of crust and lithosphere, types of volcanism, and possibly surface dynamics and tectonics. Also, it is important to mention that C/O may be still very important for understanding silicate planets as well, because it controls the redox conditions, which may affect the size and contents of light elements in the core (and therefore physical and chemical conditions), which, for example, may affect the planetary magnetic fields.

#### 4. Sources of Abundance Variations

Having established that there are considerable variations in the abundances of individual elements in stars, and most likely planets, it is worth considering the possible sources of these variations. The elemental abundances of a stellar system depend primarily on the composition of the molecular cloud from which it forms, which is primarily set by overall galactic chemical evolution, or GCE (*e.g.*, Tinsley, 1980; Timmes *et al.*, 1995; Schönrich and Binney, 2009). A molecular cloud may differ in its chemical composition from other clouds, if it were to inherit its elements in different proportions from various sources.

There are six or seven major sites of nucleosynthesis in the Galaxy. One nucleosynthesis source is spallation, or

modification of a nucleus through a collision with a high-energy cosmic ray. This is a significant production channel for a few isotopes of Li, Be, and B. In general, it is not a major source of elements (Arnett, 1996).

Much more important are stellar nucleosynthesis sources involving stars with companions. Novae, thermonuclear explosions on the surface of white dwarfs accreting matter from a companion, create products of high-temperature hydrogen burning, including isotopes of Na, Mg, and Al. Type Ia supernovae (SNeIa) involve the thermonuclear explosion and complete destruction of a white dwarf interacting with a companion. SNeIa produce alpha elements, so called because they can be produced by successive addition of alpha particles to atomic nuclei (C, O, Ne, Mg, Si, S, Ar, Ca, Ti, Cr, Mn, Fe, and Ni). SNeIa are the sources of more than half the iron in the galactic disk and produce significant amounts of Si and Ca. Neutron star–neutron star mergers are predicted to eject  $\sim 10^{-3} M_{\odot}$  of very high atomic number *r*-process (rapid neutron capture process) material (e.g., Goriely *et al.*, 2011). All these examples create elements only many hundreds of millions of years after the progenitor stars form.

Single stars also contribute significantly. In low- to intermediate-mass ( $< 8 M_{\odot}$ ) stars, deep convection dredges up material processed by nuclear burning during the stars' lifetimes, typically after hundreds of millions of years. This is the dominant source of C in the universe as well as *s*-process (slow neutron capture process) elements, roughly half the species between the iron peak and  $^{209}\text{Bi}$ . Stars above  $\sim 25 M_{\odot}$ , depending on metallicity, lose their entire hydrogen envelopes to eruptive phases as luminous blue variables and continue losing substantial mass as Wolf-Rayet stars, which are essentially bare stellar cores. Wolf-Rayet winds may contain products of hydrogen burning (He with modified abundances of Na, Mg, Al), partial He burning (high C abundance), or complete He burning (O-dominated with modified Ne abundances). Finally, stars above  $\sim 8 M_{\odot}$  can explode as core collapse supernovae (CCSNe), which are the primary production site of O, Si, and a significant fraction of Fe. CCSNe eject many isotopes of elements with both even and odd proton numbers, for example P (unlike SNeIa, which produce mainly even-*Z* elements). CCSNe also produce *r*-process elements, which are the remainder of the nuclei beyond the iron peak, including all nuclei heavier than  $^{209}\text{Bi}$ . Core collapse supernovae take place from a few to tens of millions of years after the formation of the progenitor star.

It is possible that stars *within* a cloud may differ chemically, if parts of the cloud can be contaminated by the products of stellar nucleosynthesis during the few tens of millions of years during which it is forming stars. The most natural candidates for such self-enrichment are Wolf-Rayet winds and core-collapse supernovae, since their progenitors live only  $\sim 3$  to 20 Myr before ejecting elements. Strong evidence exists for this process from isotopic analyses of meteorites in our Solar System, which show that it contained short-lived radionuclides such as  $^{41}\text{Ca}$  ( $t_{1/2}=0.15$  Myr),  $^{26}\text{Al}$  ( $t_{1/2}=0.71$  Myr), and  $^{60}\text{Fe}$  ( $t_{1/2}=2.3$  Myr) at levels well above those expected from GCE (Wadhwa *et al.*, 2007; Huss *et al.*, 2009). This points to an injection of supernova material of  $\sim 10^{-4}$  of the mass in the protoplanetary disk, less than a few  $\times 10^5$  years before the formation of the first solids (Wadhwa *et al.*, 2007). Stable isotope anomalies in meteorites paint a similar

picture. Comparing the oxygen isotope ratios of the Solar System with those of nearby molecular clouds, Young *et al.* (2011) argued that  $\sim 1\%$  of the Solar System's oxygen was injected into the Sun's cloud after its composition was determined by GCE.

Even after a stellar system forms, elements can be mixed and redistributed within the star's protoplanetary disk. Planets forming at different places within the same disk may have very different abundances, especially of volatile species like  $\text{H}_2\text{O}$  and S.

#### 4.1. Injection of supernova material into forming stellar systems (Evan Scannapieco, Steven Desch)

How supernova material could be injected into a forming solar system has been an outstanding issue for over 50 years (Cameron, 1962). Previous studies have investigated how a stable molecular cloud core could be induced to collapse as supernova material was injected into it (Cameron and Truran, 1977; Boss, 1995; Foster and Boss, 1995; Boss and Vanhala, 2000), how a star-forming complex could be contaminated by waves of propagating star formation (Gounelle *et al.*, 2009), and how supernova material could be injected into an already-formed protoplanetary disk (Chevalier, 2000; Ouellette *et al.*, 2007, 2010).

More recently, Pan *et al.* (2012) studied how a single supernova can inject material into a nearby molecular cloud. Because of their intense UV emission, very massive stars ( $> 30 M_{\odot}$ ) alter the birth environment of neighboring stars. The ionization fronts they launch move slowest through the dense regions in which stars form and could also trigger further star formation at these boundaries (see review by Hester and Desch, 2005). Massive stars explode as supernovae after  $\sim 4$  Myr (Hansen *et al.*, 2004) and can potentially contaminate the protostars at the edges of these H II regions, which are typically a few parsecs away at that time.

One of the findings by Pan *et al.* (2012) was that isotropically exploding supernova ejecta do not mix into molecular clouds; but clumpy ejecta, as seen in the Cassiopeia A supernova remnant (Fesen *et al.*, 2001; Hwang *et al.*, 2004), do mix. These clumps resemble those that are robustly predicted to form in supernovae (e.g., Kifonidis *et al.*, 2003; Ellinger *et al.*, 2012) and inferred to be the sites of supernova dust condensation (Fedkin *et al.*, 2010). Ouellette *et al.* (2010), Ellinger *et al.* (2012), and Pan *et al.* (2012) modeled these ejecta clumps (which arise interior to the H- and He-burning shells) as having masses  $\sim 10^{-4} M_{\odot}$ , diameters  $\sim 10^{16}$  cm, ejection velocities between 2000 and  $6000 \text{ km s}^{-1}$ , and densities  $\sim 4 \times 10^{-19} \text{ g cm}^{-3}$ . These densities are similar to those of a typical molecular cloud, allowing the clumps to penetrate and mix into it. It is essential to understand that such clumps are extremely heterogeneous in composition. The makeup of each clump reflects the abundances in the particular part of the star/explosion where it formed (Young *et al.*, 2009; Ellinger *et al.*, 2012). Small-scale chemical mixing of clumpy supernova material is minimal. For example, the composition of one supernova bullet that formed from the O-rich shell of the star will have a completely different composition from one formed out of the C- and Ne-burning shell of the same object.

The interaction between such a clump and a nearby molecular cloud, as modeled by Pan *et al.* (2012) using the

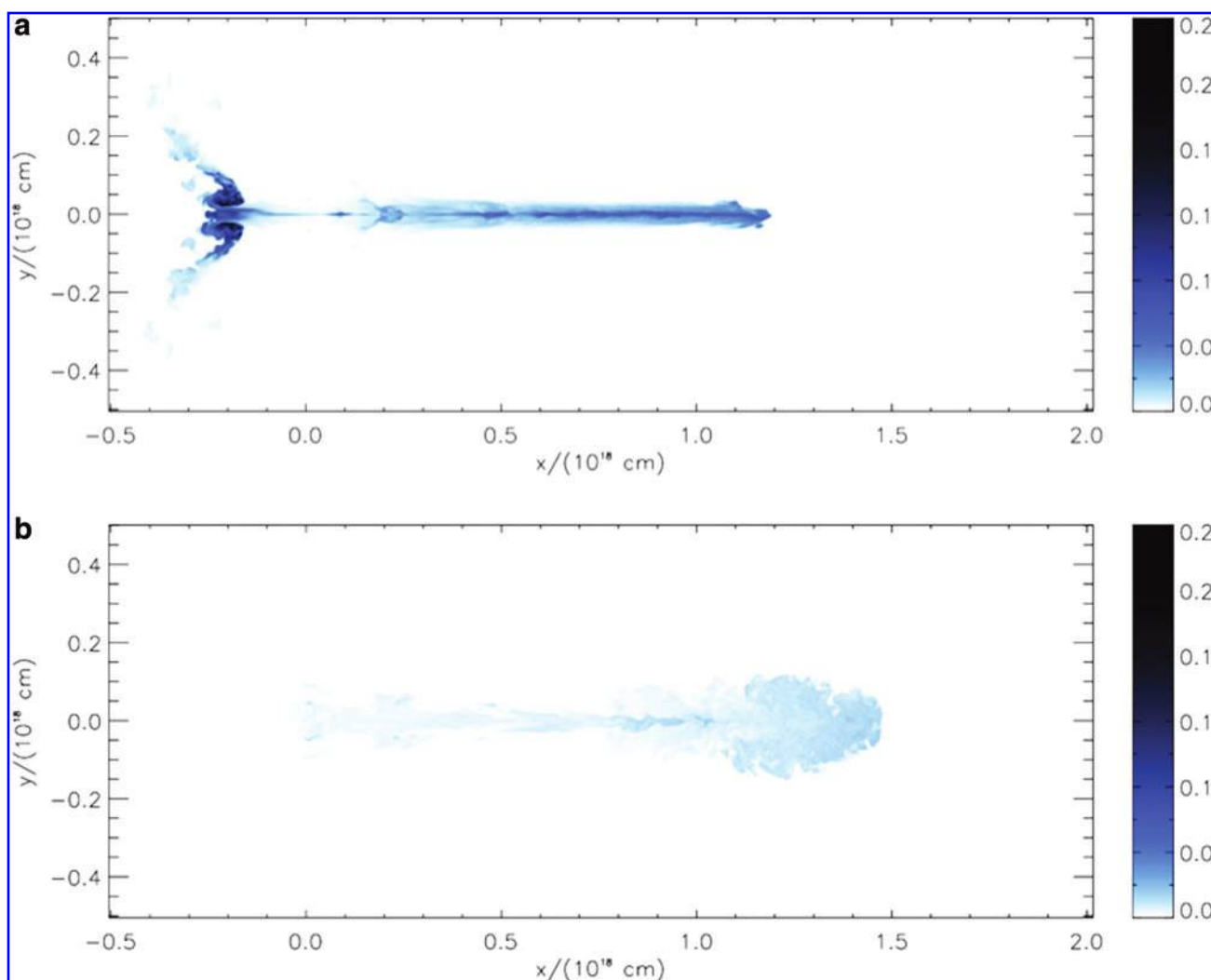
FLASH hydrodynamics code (Fryxell *et al.*, 2000), is shown in Fig. 8. Initially, the clump (a small dot on the scale of the figure) is just outside the molecular cloud at  $x=0$  and moving to the right at  $2000 \text{ km s}^{-1}$ . At  $t=5 \text{ kyr}$ , it has carved out a channel  $0.4 \text{ pc}$  long and deposited ejecta material along it, and some material has escaped back into the H II region to the left. At  $t=30 \text{ kyr}$ , the clump has stalled at a depth of  $0.5 \text{ pc}$ , and  $86\%$  of the clump mass remains in the cloud.

Subsequent turbulent mixing will distribute the ejecta laterally, so that all the gas within  $0.5 \text{ pc}$  of the edge of the H II region will be contaminated. The roughly  $2 M_{\odot}$  of supernova ejecta will thus mix with about  $4700 M_{\odot}$  (assuming the ionization fronts are  $4 \text{ pc}$  from the massive star). All stars in the act of forming at the time of the supernova will be contaminated with ejecta at a level  $\sim 10^{-4}$  to  $10^{-3}$ , which is sufficient to explain the abundance of  $^{26}\text{Al}$  and potentially other short-lived radionuclides (Pan *et al.*, 2012). Because

the clumps are  $\sim 100$  times more metal-rich than a solar composition gas, they can alter the abundance of oxygen or other elements by  $1\text{--}10\%$  as well (Pan *et al.*, 2012), possibly explaining the isotopic shifts inferred by Young *et al.* (2011). Based on the rates of triggered star formation estimated by Snider *et al.* (2009), Pan *et al.* (2012) estimated that  $10\text{--}20\%$  of all Sun-like stars may experience such  $1\text{--}10\%$  shifts in elemental abundances, especially in primary isotopes like  $^{12}\text{C}$  and  $^{16}\text{O}$ .

#### 4.2. Snow lines in externally photoevaporated disks (Steven Desch)

Numerous models predict efficient radial transport of elements in a protoplanetary disk, and redistribution of volatiles in particular. A famous example is the case of water, which can diffuse as vapor out of the inner parts of disks,

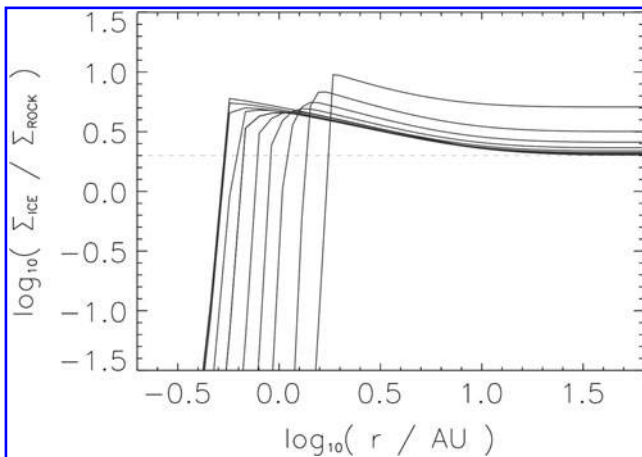


**FIG. 8.** Concentration of supernova material in a molecular cloud, following injection of a “bullet” of supernova ejecta. The ejecta clump (see text) enters the molecular cloud from the H II region to the left, at  $2000 \text{ km s}^{-1}$ , striking the boundary of the cloud at  $x=0$ . Color denotes the mass fraction of ejecta at each location. (Top) The ejecta  $5 \text{ kyr}$  after impact, showing deposition of ejecta along a narrow channel in the cloud, as well as high concentrations in a plume escaping through the channel into the H II region. (Bottom) The ejecta  $30 \text{ kyr}$  after impact, when its forward momentum is arrested, showing moderate concentrations along the channel, to a depth of  $0.5 \text{ pc}$ . Eighty-six percent of the ejecta remain in the molecular cloud. Color images available online at [www.liebertonline.com/ast](http://www.liebertonline.com/ast)



condensing and remaining trapped in the colder outer portions of the disk (Stevenson and Lunine, 1988; Cyr *et al.*, 1998). As ice particles grow, they may radially drift inward with respect to the gas, bringing water back into the inner nebula. A steady-state balance, with some water vapor in the inner disk and an enhancement of ice in the outer disk, may subsequently be established (Cuzzi and Zahnle, 2004; Ciesla and Cuzzi, 2006). The boundary between these two regions—termed the “snow line”—is taken to be where the temperature  $T \approx 170$  K, the condensation temperature of water ice at nebular pressures (*e.g.*, Lodders, 2003). Large enhancements of ice, an order-of-magnitude increase above normal cosmochemical abundances, are possible beyond the snow line, while water can be virtually absent inside the snow line.

A representative calculation by Desch (2012) is depicted in Fig. 9, which shows the ratio of ice column density to rock column density as a function of heliocentric distance, at 0.5 Myr increments, in a viscously evolving protoplanetary disk. In this toy-model calculation, the disk evolves according to the standard equations (Lynden-Bell and Pringle, 1974; Hartmann *et al.*, 1998), and the temperature of the disk depends on the rate of mass accretion through the disk,  $dM/dt$ . As  $dM/dt$  decreases over time, the temperature drops everywhere in the disk, and the distance where  $T = 170$  K moves inward. At each point, water can exist as vapor or ice, and diffusion of vapor and inward migration of large ice particles relative to the gas are accounted for in a parameterized fashion, as are the diffusion of dust and inward migration of large rocky particles. As predicted, the ice/rock ratio is enhanced over average solar nebula values beyond the  $T = 170$  K line and essentially vanishes inside it. As the



**FIG. 9.** Ratio of surface density of ice to surface density of rock, as a function of heliocentric distance, at various times in the evolution of a protoplanetary disk. Dashed line denotes the cosmochemical ratio. From right to left, the times depicted are disk ages of 0.5–5.0 Myr, incremented by 0.5 Myr. Water is in vapor form, not ice, interior to the radius where  $T = 170$  K (the snow line); ice is enhanced above solar abundances (dashed line) beyond the snow line, as vapor diffusing outward is cold-trapped there. Outward diffusion of vapor is balanced by inward migration of large icy solids. Over time, as mass accretion decreases, the temperature in the disk decreases and the snow line moves inward.

nebula cools, the region of enhanced ice/rock ratio, the snow line, moves inward as well.

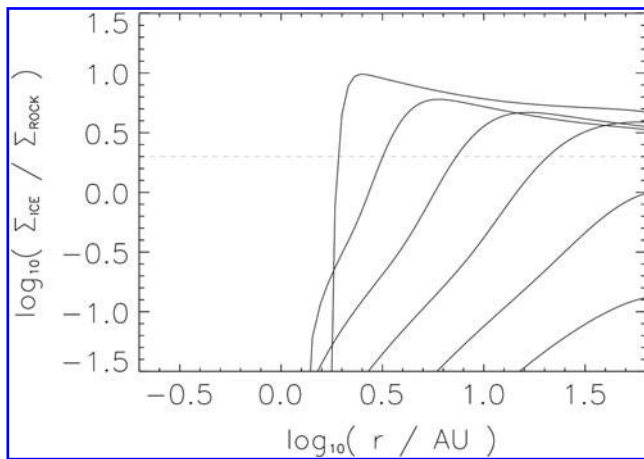
Such snow line models are invoked to explain the distribution of water within our solar system. S-type asteroids, associated with relatively dry ( $< 0.1$  wt % water) ordinary chondrites, are found predominantly inside 2.7 AU from the Sun, while C-type asteroids, associated with relatively wet ( $> 10$  wt % structurally bound water) carbonaceous chondrites, are found outside 2.7 AU. As it happens, the *blackbody temperature* of an isolated grain at 2.7 AU, exposed to the present-day Sun, is 170 K. Thus 2.7 AU is taken to be the snow line in the solar nebula. But as can be seen in Fig. 9, temperatures are cold enough to condense ice well inside 2.7 AU. Ordinary and carbonaceous chondrites formed while the disk gas was still present, at an approximate disk age of 3 Myr. Because disks are dusty and absorb sunlight obliquely, temperatures in the disk are far colder than the blackbody temperatures at the same locations. Viscous dissipation can add accretional heating as mass accretes through the disk at a rate  $dM/dt$  through each annulus, but the observed mass accretion rates in disks 3 Myr old are relatively low,  $dM/dt < 10^{-8} M_{\odot} \text{ yr}^{-1}$  (Gullbring *et al.*, 1998). Under these conditions, models of disk temperatures almost universally predict  $T < 170$  K no farther from the Sun than 2 AU (Lecar *et al.*, 2006), and more detailed models that account for the vertical structure of the viscous dissipation would put the snow line inside 1 AU (Lesniak and Desch, 2011). If a snow line existed in the solar nebula, it was probably closer to 1 AU, certainly well inside 2.7 AU. This raises the question of why the inner Solar System is so dry. Why did Earth and Mars not accrete as much ice as rock?

A possible solution to the problem, proposed by Desch (2012), is photoevaporation of the Sun’s protoplanetary disk by the FUV radiation from a nearby massive O star. Desch (2007) demonstrated that externally photoevaporated disks can experience steady-state net outward transport of mass from their inner regions to the disk edge being photoevaporated (typically at 50 AU for a disk several parsecs from an O6 star; Adams *et al.*, 2004). Desch (2012) hypothesized that this outward mass flux could prevent the inward drift of icy particles that balances the outward diffusion of water vapor across the snow line. Figure 10 depicts the same disk modeled in Fig. 9, only this time with external photoevaporation from an O6 star 2 pc away included. Vapor diffuses out of the inner solar nebula but is never replenished, and the inner Solar System can lose essentially all its water ice despite being more than cold enough for ice to condense.

## 5. Implications and Future Directions

Variations in element abundances among stars have consequences for the habitability of the planets orbiting them in ways that we are only beginning to discern but that promise the birth of entirely new multidisciplinary fields.

Historically, astrophysics and biology have intersected only at the surface of a planet, but it seems increasingly likely that the geochemistry and geophysics of planetary interiors—which ultimately shape surface environments—are also strongly affected by abundance variations of the magnitude that is observed among nominally Sun-like stars. It is possible to envision vast new research frontiers opening up as mineral physicists, petrologists, and geophysicists begin to conduct



**FIG. 10.** Same as Fig. 9, except for a protoplanetary disk that is experiencing external photoevaporation at its outer edge (around 50 AU). From left to right the curves depict disk ages of 0.5–2.5 Myr, incremented every 0.5 Myr. As the disk evolves, external photoevaporation causes the net flux of matter to be outward, which prevents the inward migration of solids that would replenish the inner Solar System. In only  $\sim 10^6$  years the inner solar nebula is essentially devoid of ice.

experiments and construct computational models that deviate from terrestrial compositions, taking astrophysical data as inputs and providing as outputs geophysical and geochemical constraints that will be crucial to predicting surface habitability. We propose the term “astrobiological stoichiometry” for the study of chemical compositions and their effects on star and planet formation, and planetary interior, surface, and atmospheric processes that influence the habitability of extrasolar planetary systems.

While the current research trends are exciting and tantalizing, there is a danger of overgeneralization and over-application of early results. There are critical areas where fundamental astrophysical work is needed before emerging notions of planetary habitability can mature.

First, and most obvious, is the observational determination of chemical compositions of both stars and planets. Determining accurate compositions is not straightforward, even for the most easily available targets, the host stars of extrasolar planets. Researchers in this field must perform detailed comparisons of their methods to identify the sources of discrepancies. That information must be used to agree upon a set of best practices for abundance determination, including optimal sets of atomic constants, line lists, line flux measurement strategies, and stellar parameter determination. Independent verification of stellar parameters by methods such as asteroseismology should be used where possible. Finally, it is essential to determine the abundances of as many elements as possible for any interesting host star. The Fe abundance alone is not sufficient even to accurately describe the host star’s evolution.

Second, greater progress must be made in deriving abundances of planets. For now observations are restricted to absorption or reflection spectra from jovian planets that are transiting or directly observable. The few existing observations to date already confirm that planets differ compositionally from their stars. These differences could arise from disk processes during planet formation such as volatile transport and loss from the protoplanetary disk; con-

densation of minerals; or radial mixing of gas, dust, and planetesimals. Chemical abundances can be changed as a planet accretes, and elements can partition differently within the planet itself. This is illustrated by the apparent lack of oxygen in Jupiter’s atmosphere as probed by the Galileo probe, which Juno may confirm to be a global depletion of oxygen from Jupiter’s upper atmosphere, possibly due to dissolution of oxygen with neon into helium droplets. Terrestrial exoplanets so far cannot yield any direct composition information, only estimates of bulk density when their radii are known from transits and mass from orbital elements.

While observations of stars and planets improve, other observations can help constrain how stellar compositions translate into planetary abundances. Dust in young systems’ debris disks left over from planet formation, produced by collisions between small parent bodies, represents material as it condensed out of a protoplanetary disk but before it was substantially processed in full-fledged planets. Near IR and mid-IR spectroscopy of such disks can be used to identify mineral species, giving insight into elemental ratios, potentially even as a function of distance from the star. A second source of dust, seen in a handful of systems, is impacts between large planetary bodies. Astromineralogical studies of these systems can provide insight into the mineral species that formed in high temperatures and pressures in planetary interiors.

A goal for the future is to model the creation of planets from an initial nebular composition to a planetary body, with the caveat that many steps along the way are stochastic. Dust condensation models exist for non-solar compositions and can be generalized further (Ebel, 2000, 2013). Disk models can compute how volatiles and minerals are transported through the disk and provide initial conditions for models of planetesimal and planet assembly. Migration of planets during and following assembly can be included. While the stochasticity of all these processes requires a large number of realizations of simulations, stellar compositions should be “mappable” onto planets in an average sense. After formation of a terrestrial planet, various planetary processes are open to exploration. Differentiation, thermal evolution, convection, mineralogy and rheology, and exchange between atmospheres and planetary surfaces through outgassing, erosion, and sequestration are all active fields of research. This set of problems will hopefully serve as an impetus to increase existing interaction between these fields and initiate new connections.

The fundamental purpose of astrobiology is to discover life on a world around another star. As our search has progressed, the options have multiplied: recent estimates are that 20% of Sun-like stars harbor a terrestrial planet at orbital distances that would yield Earth-like habitability. For the first time, astrobiology is faced with the problem of narrowing the parameters and creating a refined target list in the search for life. Astrobiological stoichiometry provides a robust strategy for hunting for life. As the field develops along the outlines explored here, we will be able to use readily observable stellar and planetary properties, especially chemical composition, and make predictions about the surface conditions on planets that are candidates for detailed follow-up in a next generation of missions searching for biosignatures. Compositional information is a key component

in predicting the long-term climate evolution and surface geochemistry of extrasolar planets. Surface geochemistry, in turn, determines the total free energy available to an ecosystem and makes possible comparisons between the rates of abiotic chemical cycles and metabolic rates for hypothetical biological communities. These comparisons are vital to predicting the likelihood of a planet having detectable biosignatures. Chemical abundances are nearly as fundamental to habitability as stellar types and planetary orbits.

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No competing financial interests exist.

### Abbreviations

CCSNe, core collapse supernovae; EOS, equation of state; FUV, far ultraviolet; GCE, galactic chemical evolution; HZ, habitable zone; SNeIa, type Ia supernovae.

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